Paleoseismology, methods and examples

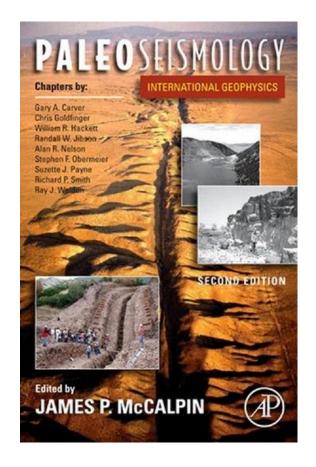
Paleoseismology

- behaving of seismogenic fault in geological history

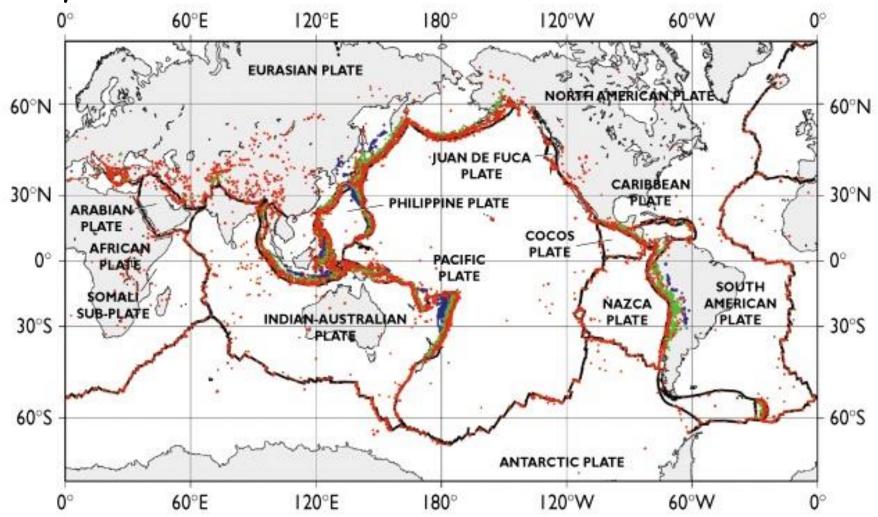
Paleoseismology studies **<u>prehistoric</u>** earthquakes from geological record

Seismologists - data measured instrumentally during earthquakes

X Paleoseismologists interpret **geological phenomena** accompanied by individual EQs



Why?



Present day seismicity - plate boundaries, intraplate regions Catastrophic EQs - sometimes in areas with faults with no present day seismicity, seismic cycle - longer reccurence interval (China, New Zealand) Most areas - record of historical EQs only several hundred yrs (historical and instrumental seismicity)

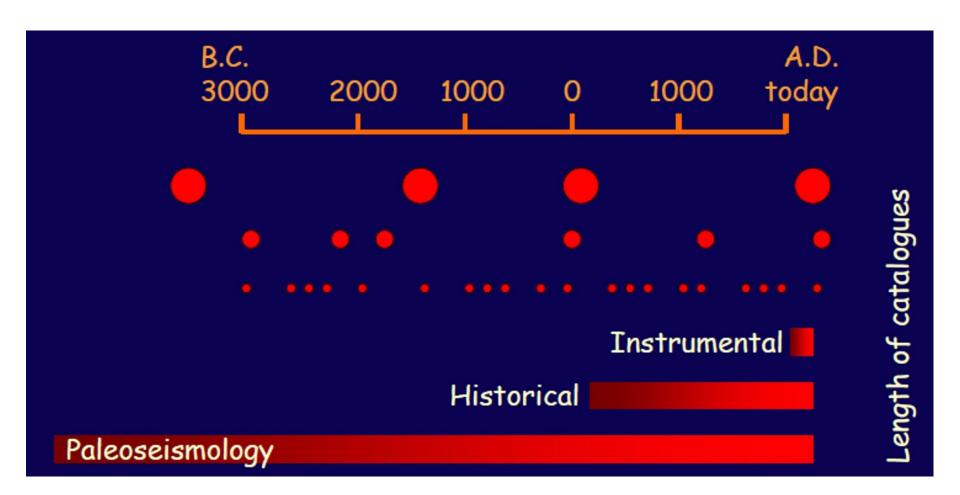
X some active faults expressed in morphology and geology – no historical seismicity or large EQs

China and Middle-East - record thousands yrs and more, still not long enough; fault active millions yrs - 3,000 yrs - only little part of faulting history

Seismic hazard assessment - based on very short period of record of historical EQs, it may cause 2 problems:

 overestimation of probability of future EQs based on historical large EQs, but with long recurrence interval (seismic energy is released)

 underestimation - in areas with seismogenic faults but no historical record (strain accumulation)





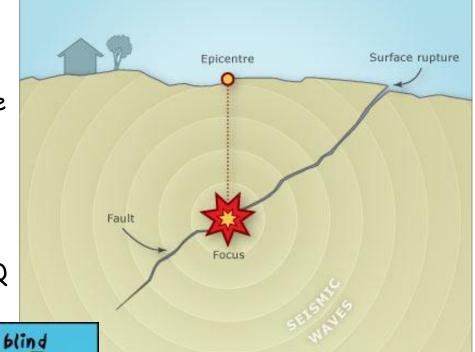
Paleoseismology extends record of EQs into the geological past

Earthquakes catalogues too short

Premise - EQ only larger M > 6 can create permanent deformation on the surface \rightarrow topografic instability \rightarrow new processes erosion and accumulation \rightarrow new landforms and sturctures \rightarrow geological record of EQ

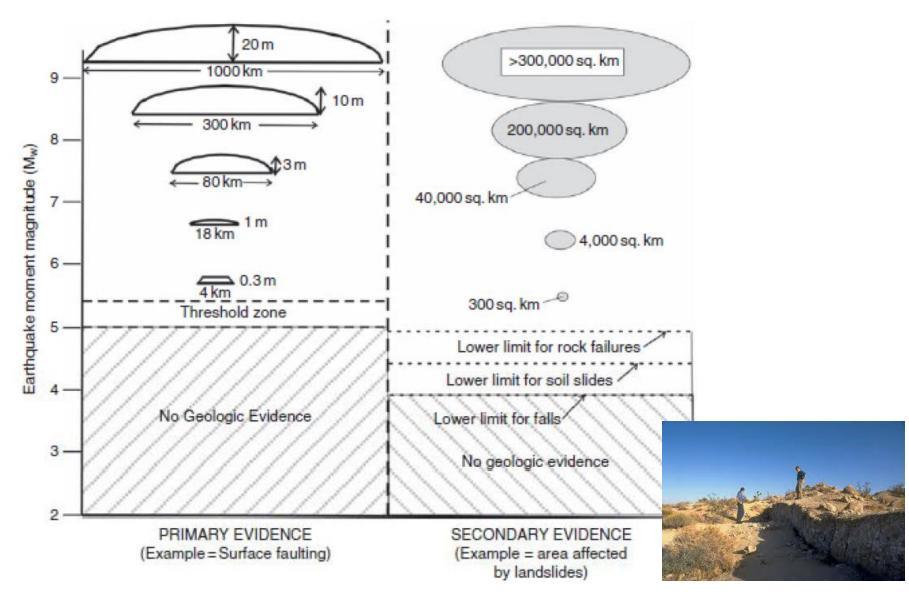
hidden

surface breaking



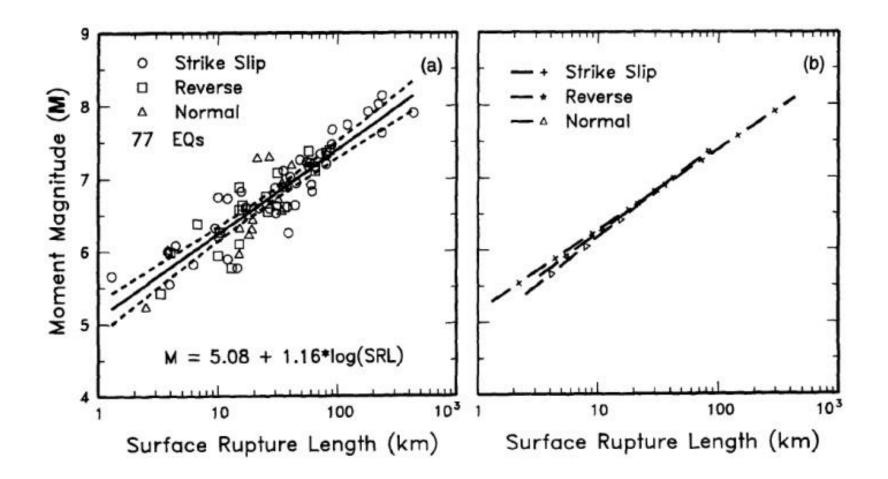
Smaller EQ - rarely geological expression created or survives Fault type - normal faults $M \ge 6.3$; strike-slips - California i - M = 6.25-6.5, Depth of seismogenic crust - deeper needs higher magnitude

> Loma Prieta 1989 M=6.9, 2m slip in depth 3-18km, no surface rupture Gujarat 2001 M=7.7, blind fault, 1-4m in depth 9-15km,

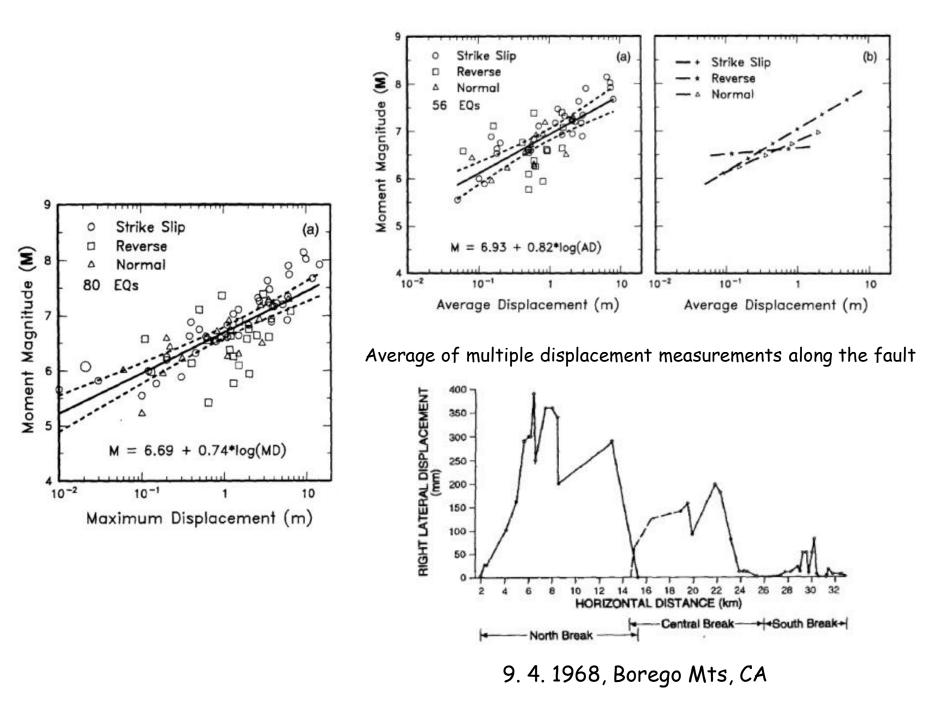


Empirical relationships based on observation from historical EQs

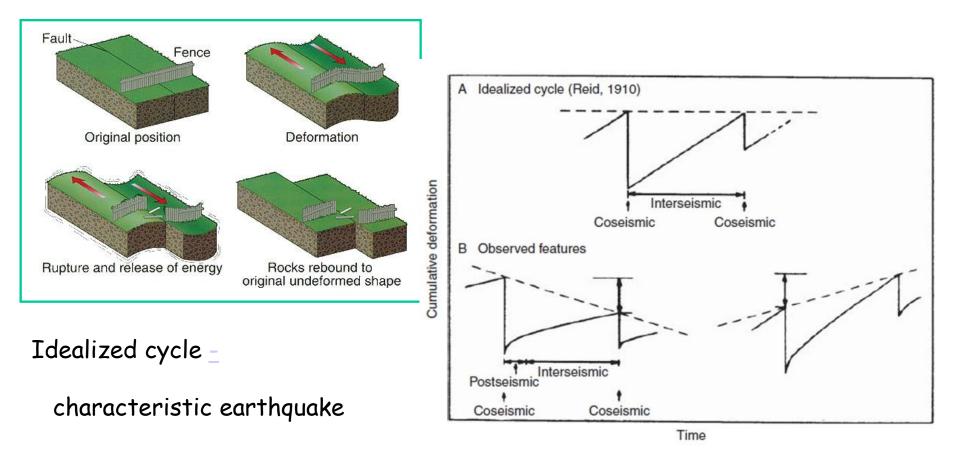
Relationships: fault length, amount of displacements, size of Magnitude e.g. fault 80km long can generate EQ Mw=7.5 and displacement 3m



Empirical relationships - historical EQs (421), focis depth <40km, Mw > 4.5 Wells, and Coppersmith 1992



Seismic cycle – elastic model



Earthquake deformation cycle

Paleoseismological study of faults

- Localisation and geometry (geomorphology, geological mapping)
- Slip rate faulting velocity (= displacement/time)
- Slip per event characteristic displacement during individual EQs
- Recurrence period (repeated EQ, frequency EQ)
- Elapsed time time from the last EQ
- Maximum potential magnitude

Chronological reconstruction of movements

 stratigraphic, structural, geomorphological, biological, archeological evidence

* dating of displaced features or movement indicators





 dating of multiple movements (EQs) - recurrence interval, long-term sliprate, vaiability of movements during EQs



predict localisation and magnitude of future EQs

Methods

- direct observations of dislocated objects - on the surface or in **trenches**, outcrops



* young sediments, fine grained, stratified - well recognizable displacemnt of layers, not thick

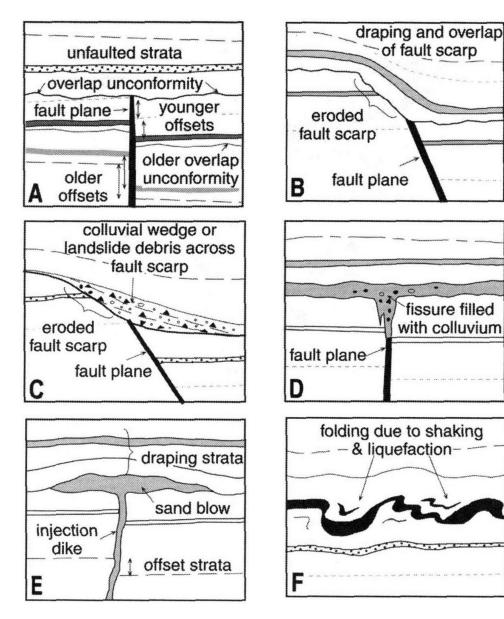
Alluvial fans, lake sediments X debris flow

 \ast datable material- chronology of movements

Evidence of earthquakes (EQ) in geological profiles in a trench

- A) Difference in cumulative offset
- B) Buried fault scarp
- C) Coluvial wedge- typical for sudden movement
- D) Filled fissures by overlying material
- E) Sand dykes

F) Liquefied layers



Allen (1986)

Repated EQs

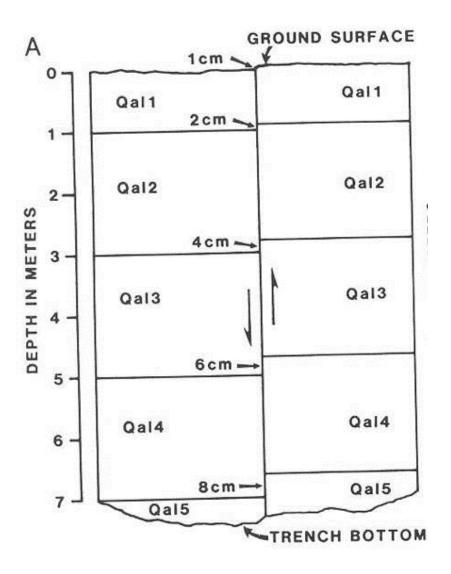
- Difference in cumulative offset
- ? How many: retrodeformation

4 events - vertical offset 2cm

Oldest layer - (Qal5) all 4 events, cumulative 8cm

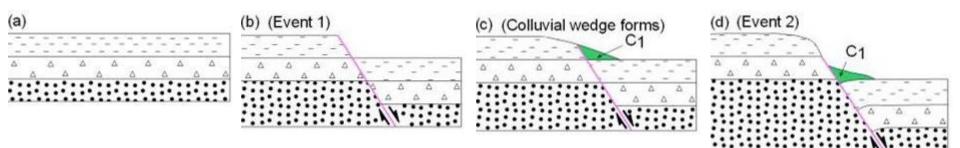
Youngest (Qal1) has experienced only 1 event → 2 cm on the layer base, but 1 cm on the surface!

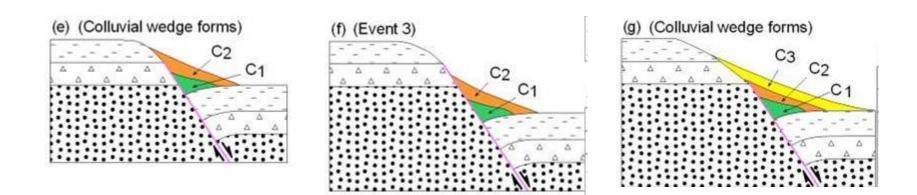
Surficial erosion



Normal faulting

Colluvial wedge

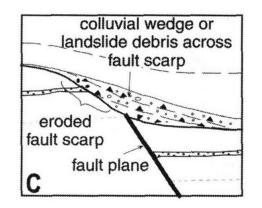






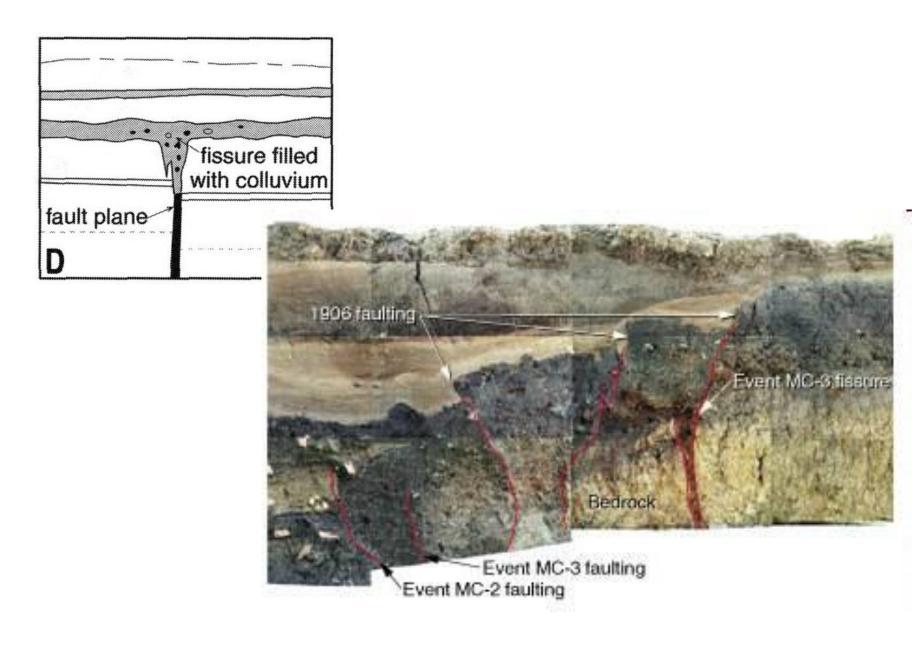
Gravitational instabilty

Fault scarp derived material - wedge









Aremogna-Cinquemiglia fault - Italy

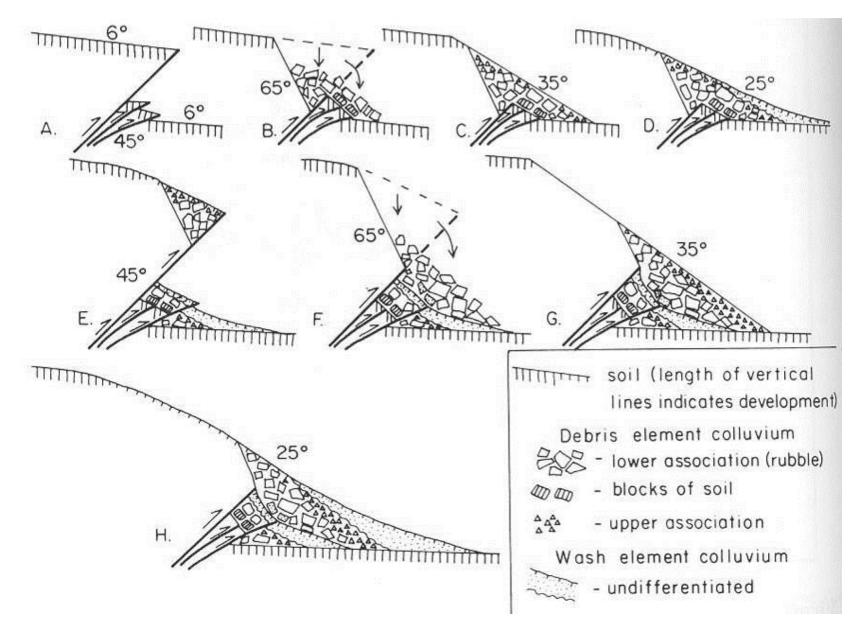


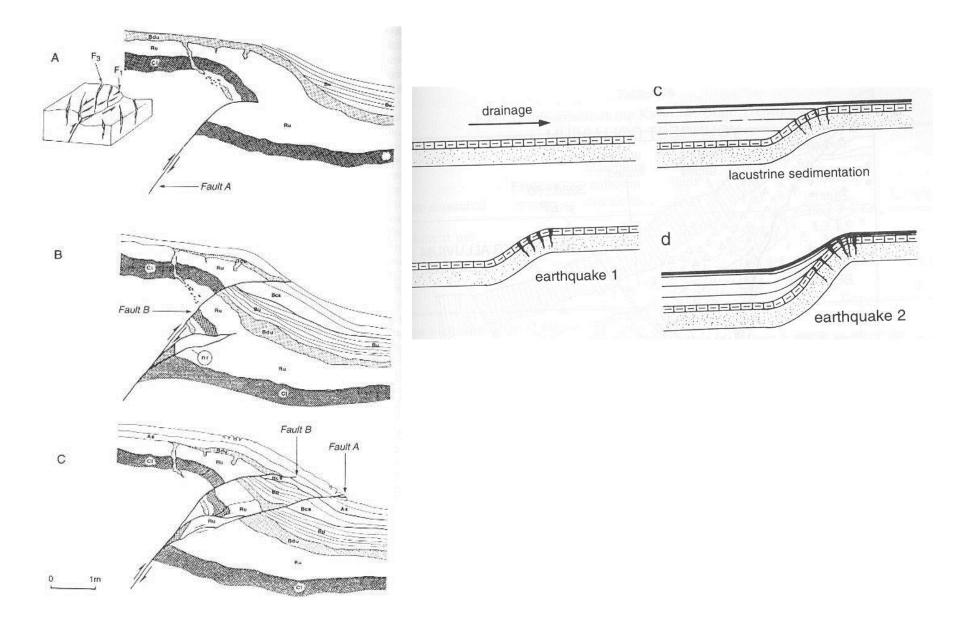


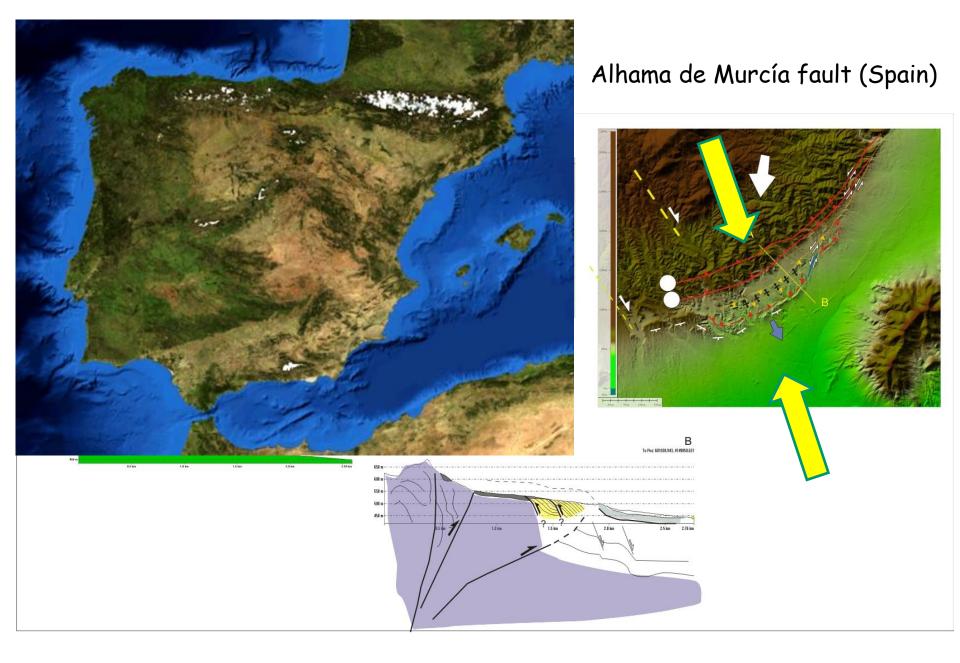


Kyrgyzstan

Reverse faults - colluvial wedge

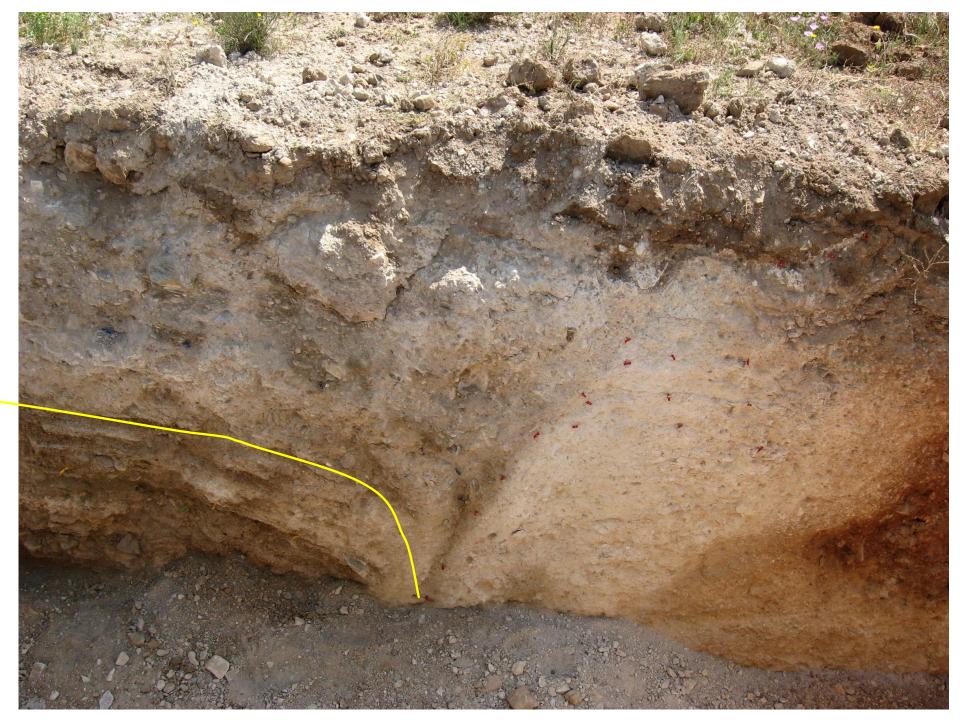




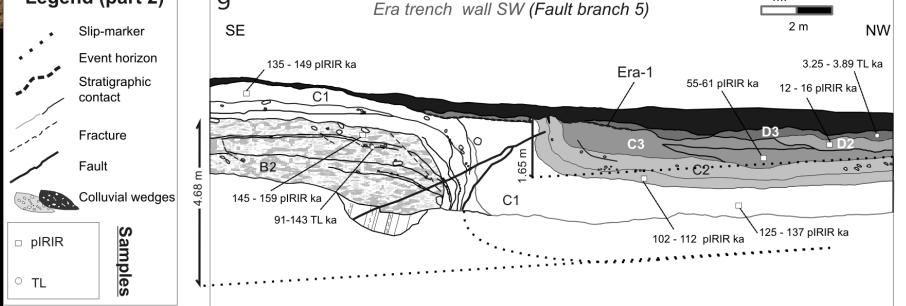


Various kinematics related to different stress diection





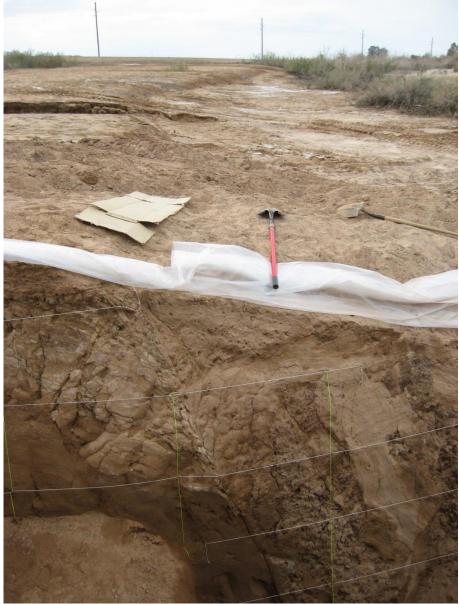




Fault scarp and colluvial wedge on strike-slip fault



Imperial fault, 1940 M=7.6m offset, 60km

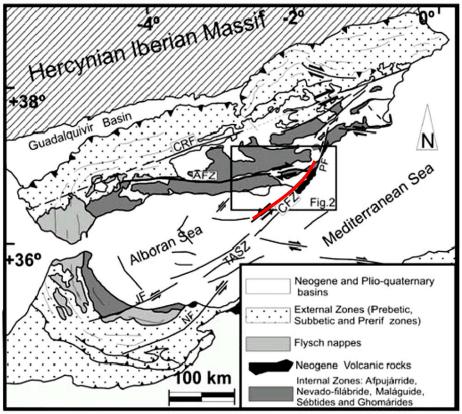




Case study

Carboneras fault zone - Spain

Carboneras





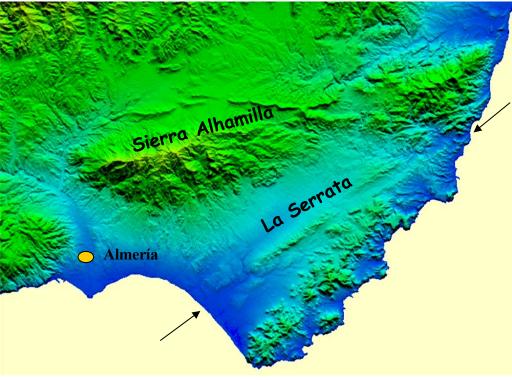
- Collision zone Europe x Africa plate
- Southern margin of Alpine orogen
- Part of Bettic Cordillera

 outer zone (nappe from Mesozoic to Tertiary rocks) paleo-margin of Iberian plate

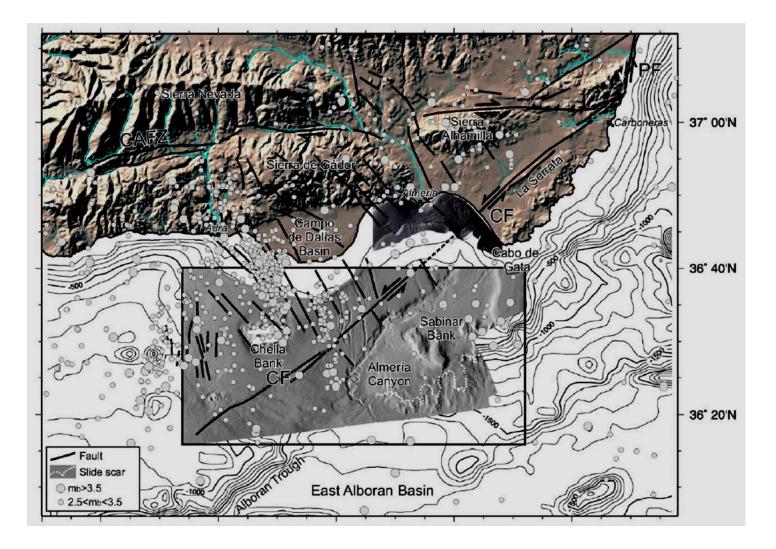
inner zone- metamorphic complex +
 Neogene to Quaternary sediments intramontane basins bounded by
 faults NE-SW

Masana E. et al.





- Carboneras formed in the last period of collision of inner and outer zone of Betic cordillera in early Miocene
- * Miocene to Quaternary stress field rotation
 - normal faults mid-Miocene part of rifting (volcanism)
 - reverse faults early Pleistocene (formation of small mountains
 - e.g. La Serrata)
 - strike-slips left-lateral (up to present-day)

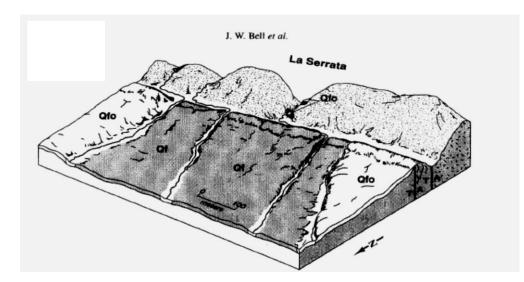


 seismicity - SE margin of Iberian peninsula - permanent shallow earthquakes M < 5,5 (transversal faults now without seismicity -Carboneras)

* last 2.000 years - at least 50 larger earthquakes

Previous studies in 90th

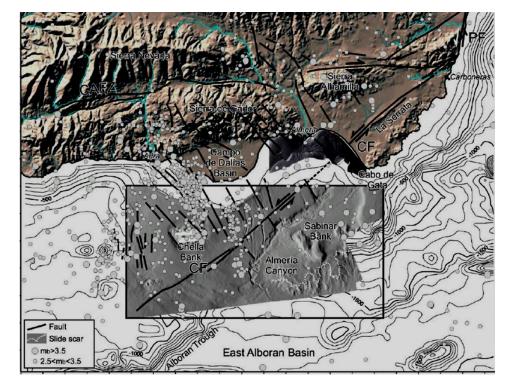
- Study of evidence of left-lateral movements dated by radiometric dating of marine terraces and their recent uplift
- 2) Measuring and dating of left-lateral movements based on offset channels

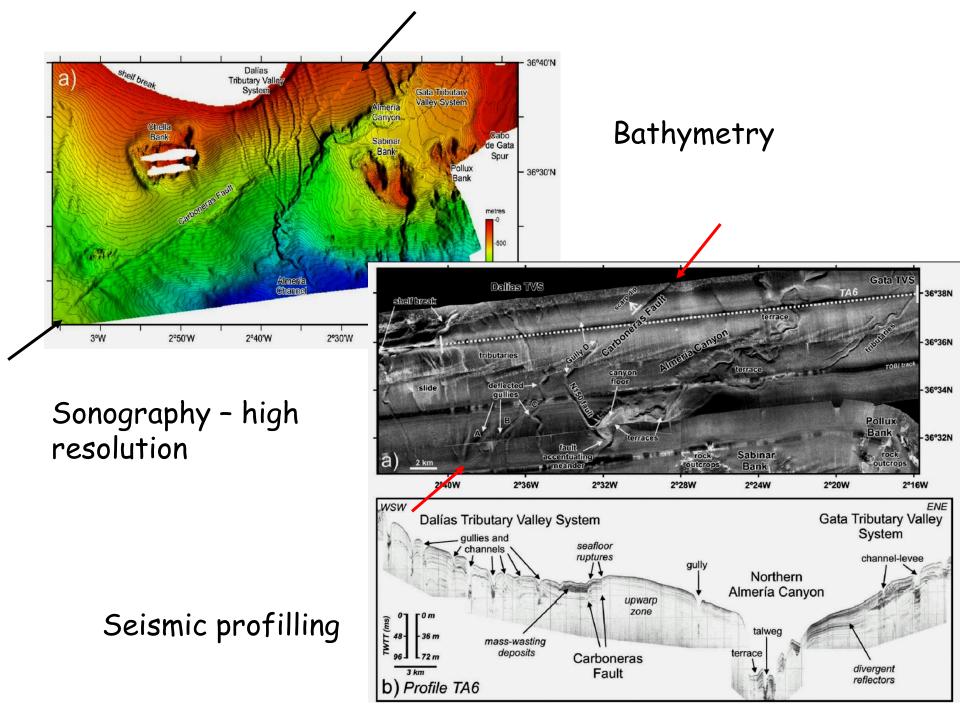


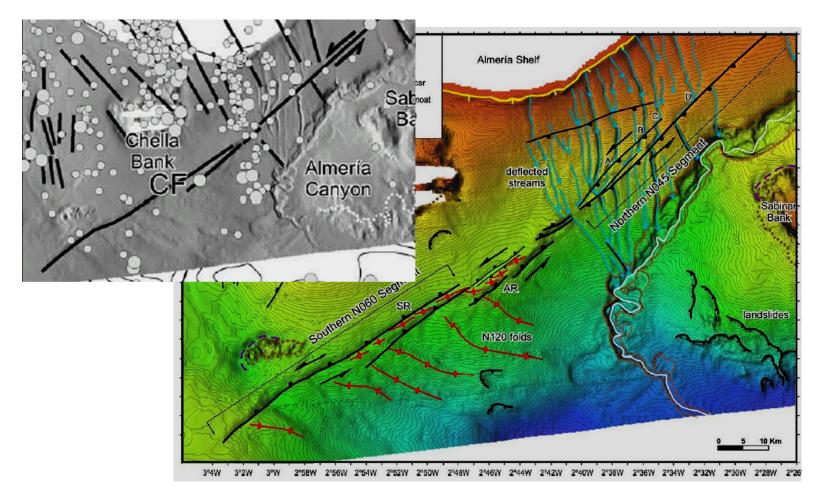
Movements in late Quaternary- relatively slow, **mainly vertical**, horizontal movements of 80-100m offset channels in La Serrata – older than 100.000 years

Methods of study of Carnoboneras fault on the sea

- Bathymetry
- Sidescan sonography
- High resolution seismic reflection
- Marinne sediments samples analysis
- * Dating of the sediments







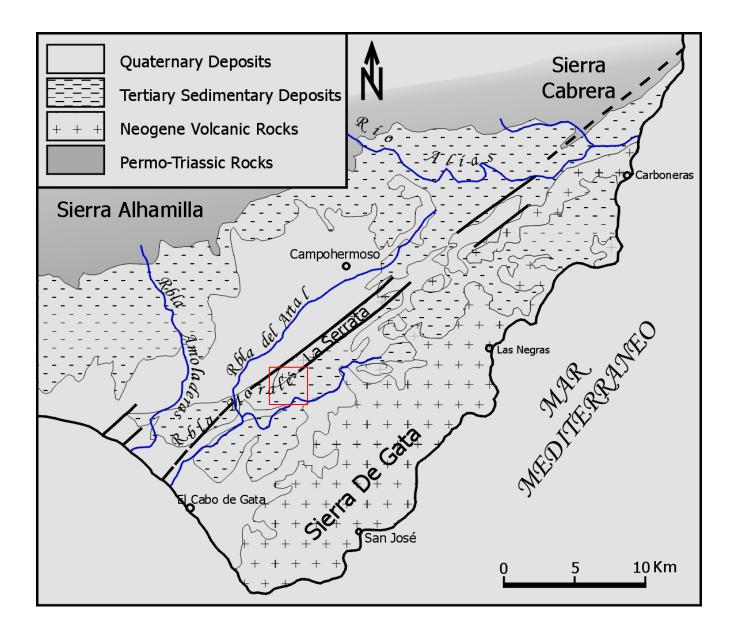
Results: Carboneras fault zone - 5-10 km wide, 100 km long, subvertical faults,

Drainage network on the inland margin - deflected,

Morphology formed by horizontal movements – pressure ridges, water gaps, late Holocene sediments, landslides.

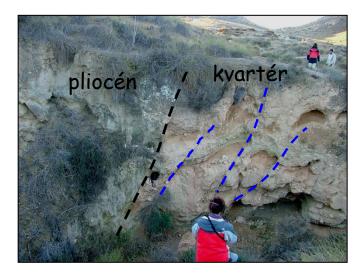
Methods of fault study on the land - offshore

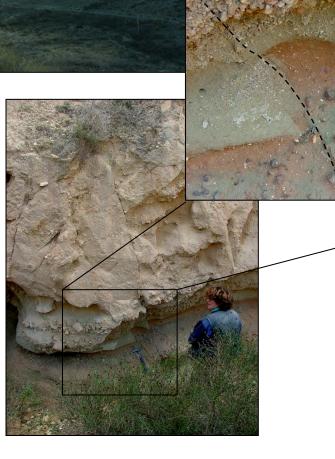
- Photointerpretation air photos
- Geomorphological mapping of dislocated landforms
- Structural mapping (faults)
- * Sedimentology (identification of generations of alluvial fans)
- Microtopography (total station)
- Geophysics (georadar, electrotomography fault tracing and goundwater level)
- Paleoseismic trenching
- * Dating of materials cut by the fault

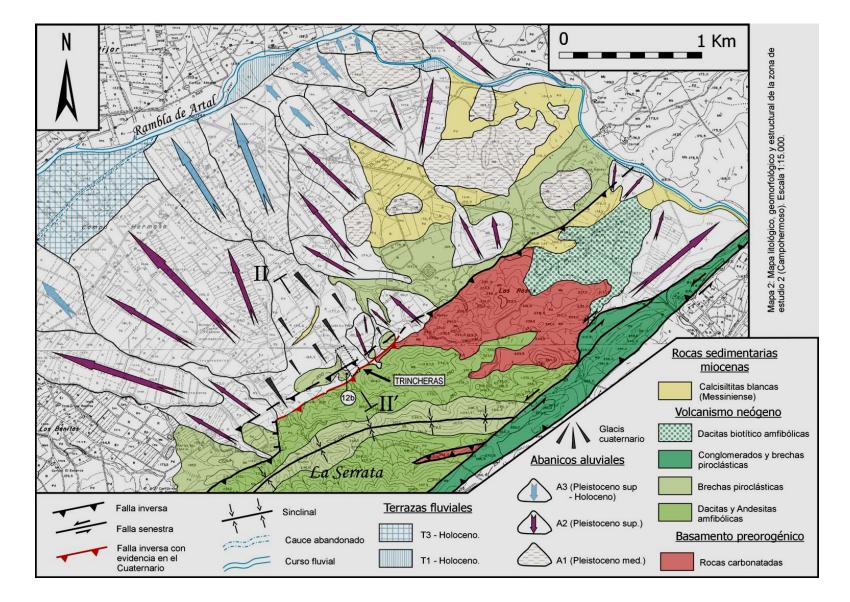




La Serrata

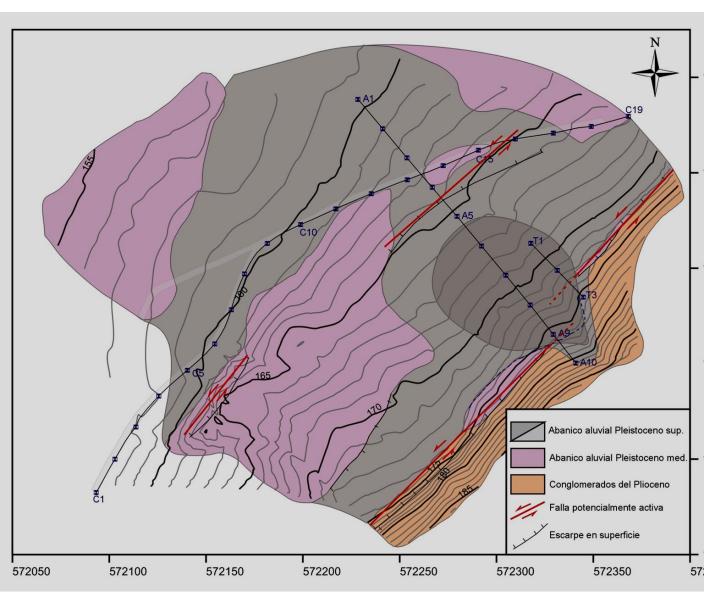






3 generations of alluvial fans - Mid and Late Pleistocene/Holocene - 3 various generations of fault movements (erosion - accummulation)

Paleoseismic trenches



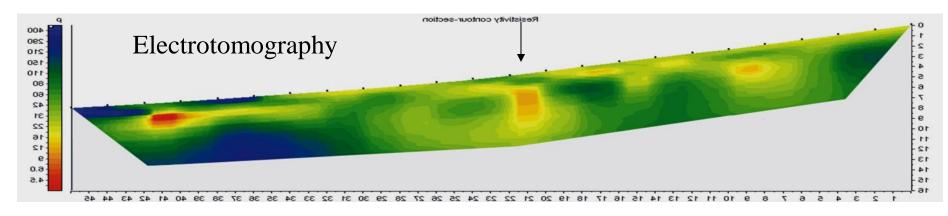
* all 3 alluvialfan generations(chronology)

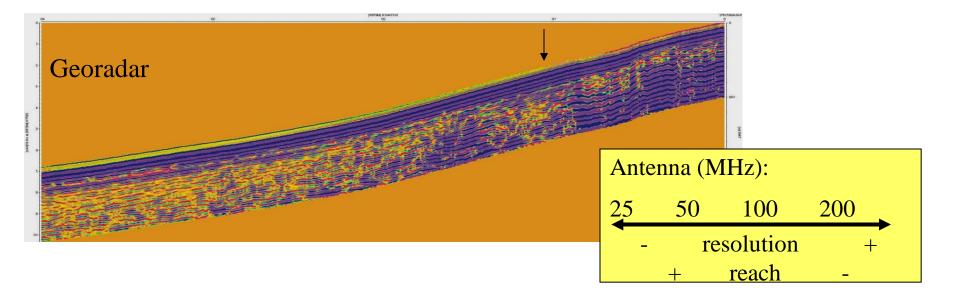
visiblemorphologicalscarps (0.7m)



Geophysics

- fault position and characterictsics of the material at the depth







Cleaning the wall, grid









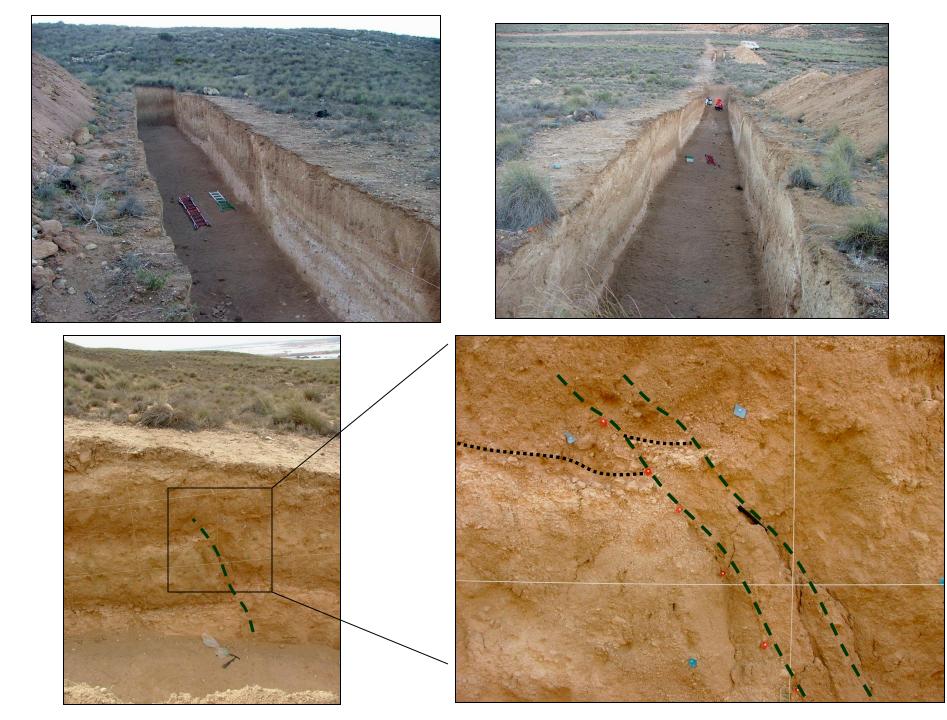
Identification of sedimentary layers and dislocations/faults











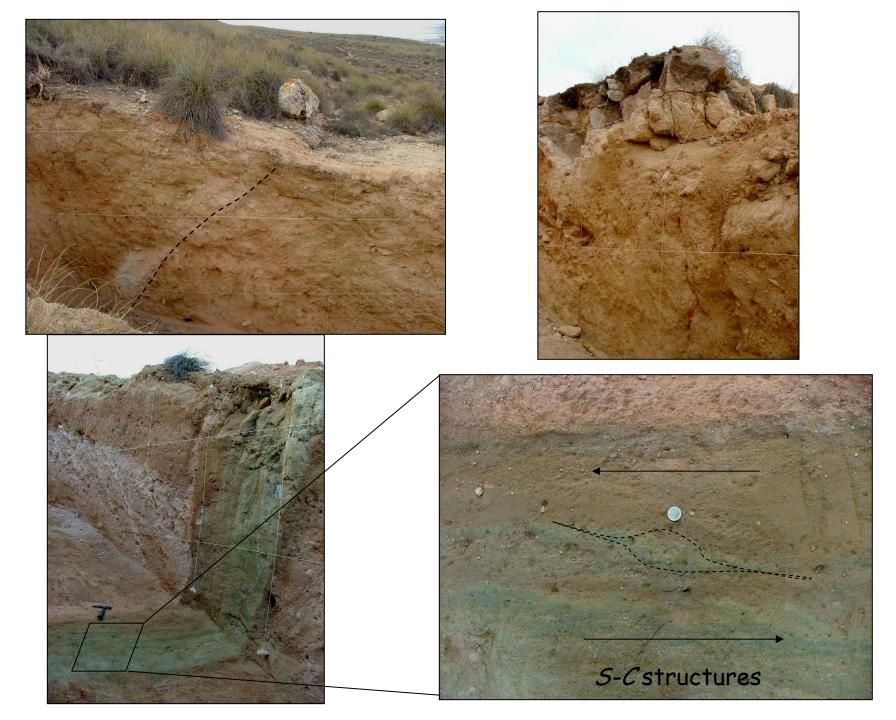


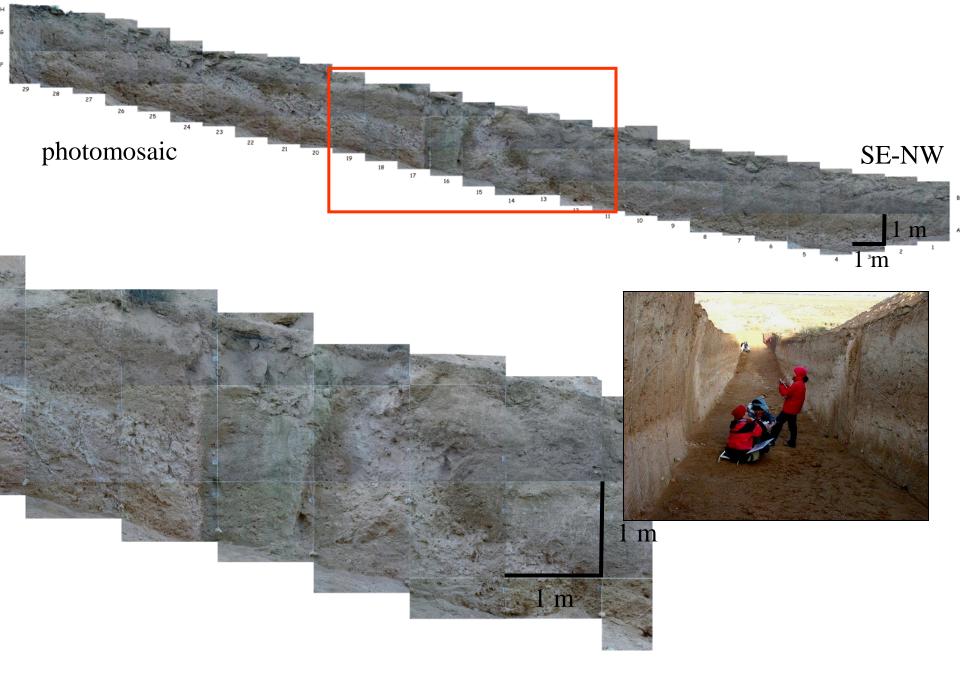
- Complex structure

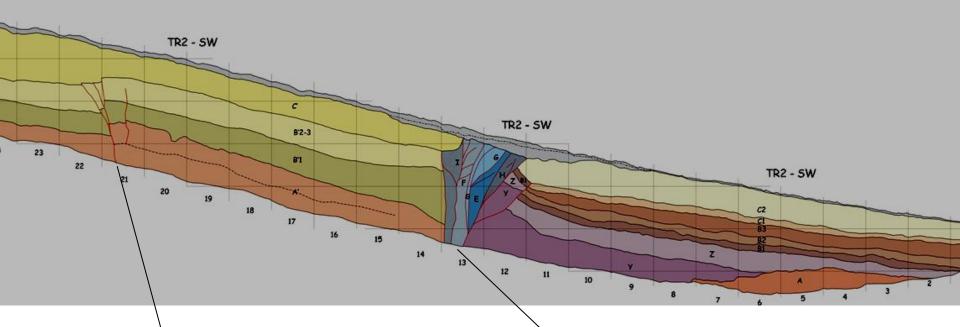
 flower structure
 transpressive regime

 Horizontal movements
 strike-slips with
 vertical
 - component-
 - Repeated movements

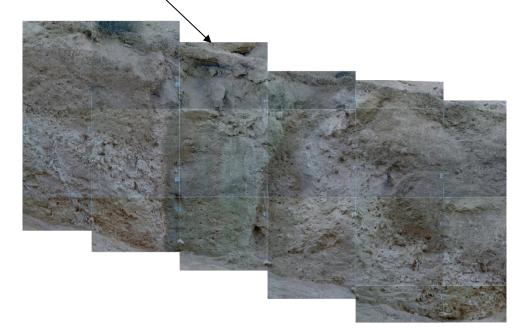




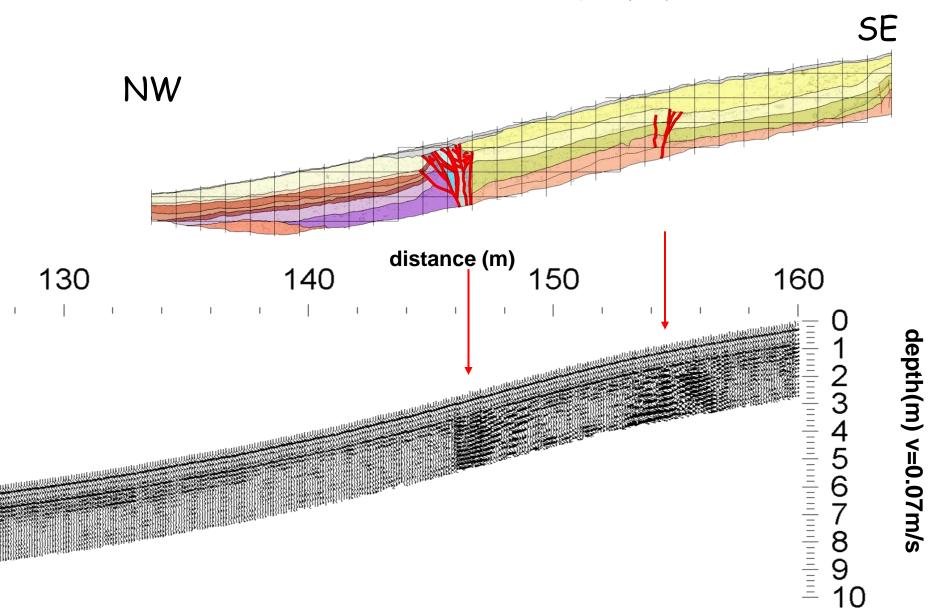








situation in situ versus geophysics

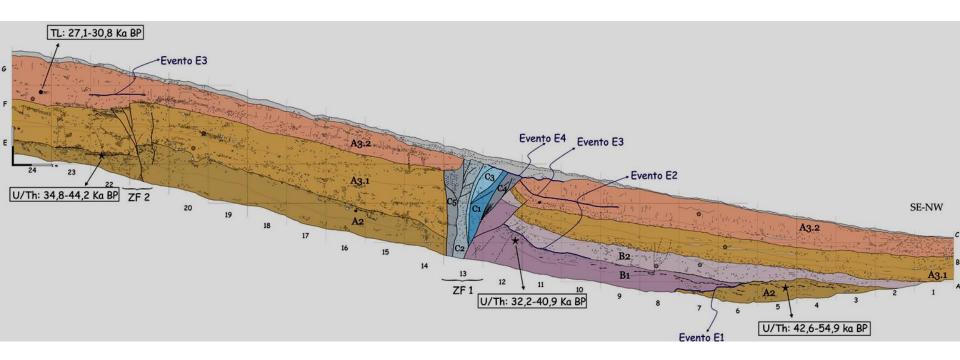


Dating – material cut by the seismic event

- * ¹⁴C radiocarbon dating method → organic material and carbonatic shells (reach 40 thousands year) – charcoal, gastropodus, organic material, wood, etc.
 - 14C in living organism, added from the environment, it decays, after death of organism, ratio of 14C/stabile 12C changes
- Optically stimulated luminiscence OSL electrones trapped in crystal lattice of sand grains - released by light activation or stimulation. (reset - zero signal). After finishing of sedimentation - signal increases due to radioactive decay. Luminiscence relased by light activation in the lab is proportional to sediments age (the time until which the electrones were accummulated - until next reset (reach 250-300 thousands yrs)
- * Thermoluminiscence $TL \rightarrow$ fine-grained sediments (100 thousands yrs)
- V/Th → carbonatic material (reach 300 thousands years) –
 laminar caliche

interpretation of trench logs, assessment of type and amount of movements

reconstruction of deformation (retrodeformation)



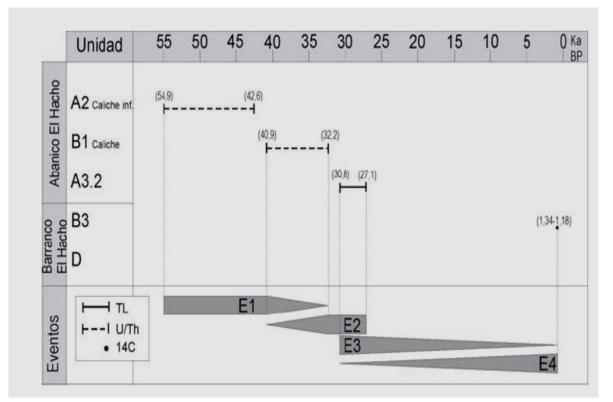
* laboratory results of dating

chronology of tectonic activity on the Carboneras fault

Results: Trench analysis and dating on Carboneras fault:

- B₁, B₂ colluvial wedge were recognized (surface degradation after suddent event) - earthquake
- Minimum 4 earthquakes during last 50 thousands years
- Recurrence period min. 14 thousands yrs
- Last event minim 1310 years ago
- Empiric relationship Magnitude X Displacement for 2 events minim.

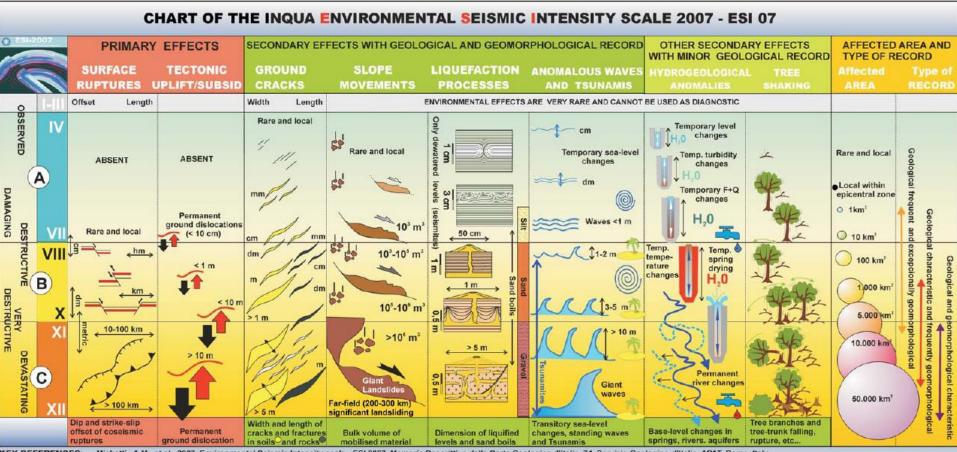
M= 6.59 and 6.97











KEY REFERENCES: Michetti, A.M., et al., 2007. Environmental Seismic Intensity scale - ESI 2007. Memorie Descrittive della Carta Geologica d'Italia, 74. Servizio Geologico d'Italia, APAT, Rome, Italy Silva, P.G., et al., 2008. Catalogue of the geological and environmental effects of earthquakes in Spain in the ESI-2007 Macroseismic scale. Geotemas, 10, 1063 - 1066, SGE, Spain Reicherter, K., Michetti, A.M., Silva, P.G., 2009. Palaeoseismology: Historical and Prehistorical Records of Earthquake Ground Effects for Seismic Hazard Assessment. Geol. Soc. London, Spec. Pub., 316 1-10. London, U.K.