

INDUCED SEISMICITY AT MIROVO SALT DEPOSIT, BULGARIA

I. Paskaleva, G. Manev, M. Kuteva

Bulgarian Academy of Sciences, Central Laboratory for Seismic
Mechanics and Earthquake Engineering, 1113 Sofia, Acad. G.
Bonchev str., block No 3, Bulgaria

Introduction

In the process of modern technologies development the social vulnerability increases because of human actions, directed in looking for new energy sources, communication systems and using of complex technological systems. These actions very often provoke induced seismicity, even in regions where seismic activity has not been observed in the past. Usually these are industrial zones where there is coal, salt or other deposit extraction processes or possible big soil mass collapse (dams, reservoirs, etc.) Similar example is the Provadia region, the only salt deposit in Bulgaria that is explored. The purpose of this study is to present some basic results of longterm observations of the dynamic processes in the region of the salt deposit.

The salt deposit

The salt diapir is claimed to be unique geological feature, several kilometers across at its surface exposure, extending and broadening to depths of perhaps 3500m. The shape of the diapir is a frustum of a cone. Its upper part is from 12-20 meters under the ground. In depth it reaches 3500m-fig.1. There is formed a salt layer in the range of depths 3500 -4500m. The deposit is imbeded in cretaceous limestones and dolomites and paleogene marlstones and it is covered by quaternary sediments. The mineral composition of the salt rock mass building the salt body is not homogenous and hallite is the predominated in this composition. The statistical analysis of the available information shows significant and irregular unhomogeneity of the rock salt massif. Rock salt physical and mechanical

Vertical cross sections of the salt diapir

1a - North-South direction

1b - East-West direction

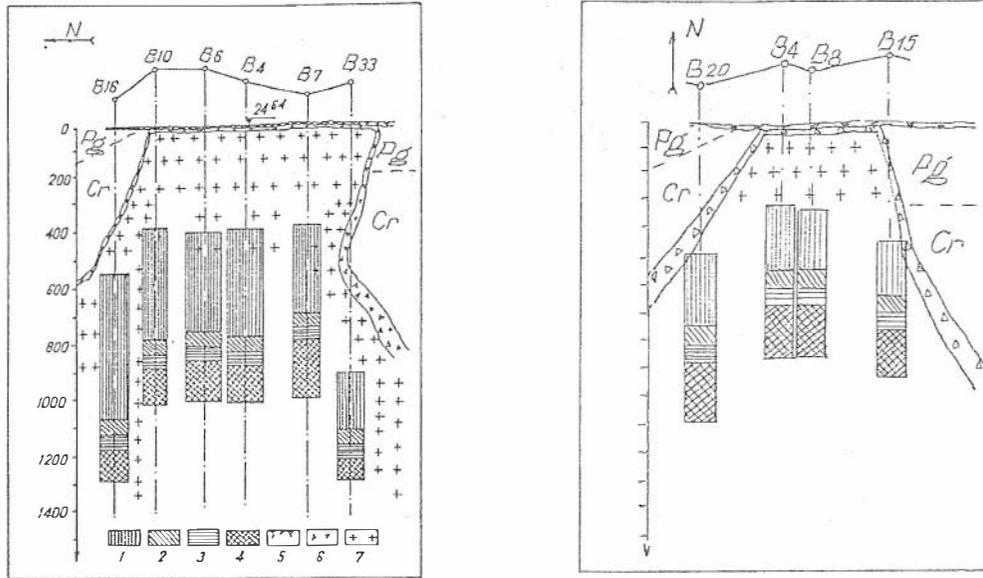
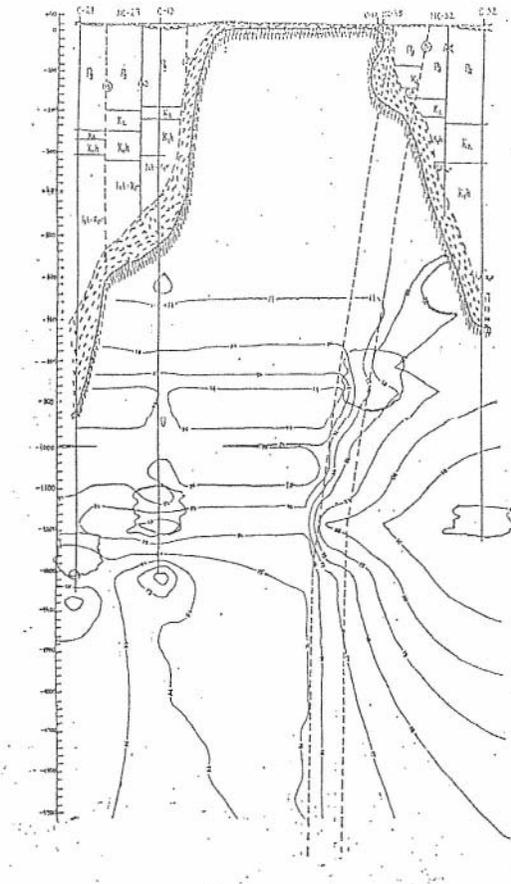


Figure 1

1 - 6 roof chamber level: 1-designed level, 2-1995 year, 3-1990 year, 4-1985 year; 5 - soil layer, 6 - breccia, 7 - rock salt.

Compressive strength
distribution in depth

Figure 2



properties are determined using test cores from more than 30 boreholes at different depths. The average in-situ density is as follows: salt-2.17g/sm.cub, insoluble dirt -2.74g/sm.cub. and salt - marl rock - 2.23g/sm.cub. The strength parameters vary in a wide range. The basic one, the compressive strength is from 8.5 to 30MPa with average values $R_c = 14 - 16\text{MPa}$ (lab test) and $R_c \text{ mass} = 5.4 - 5.8\text{MPa}$ (massif). The compressive strength trend analysis has shown significant variety in its distribution - fig.2. The rock mass rating according Bieniawski is in the range of $14 \leq RMR \leq 41$ which determines the salt diapir as a weak or very weak massif. Salt-marl rock test cores have been tested under uniaxial and threeaxial compressive longterm loading (703 hours). Quite strong expressed rheological characteristics of the salt rock have been observed. In the case of loading, overcoming with 50 percents the compressive strength, the longterm strength is evaluated in the range of 5.2 to 5.4MPa and the observed creep velocity is quite high. The stress state of the undisturbed by the faults salt rock massif at different depths and of the allerolites and argillites of the underlying formation is determined by five independent methods [5] using the so called 'rock memory'. The stress tensor variation in the salt diapir and the underlying formation is shown on fig.3. Horizontal tectonical ground movements are registered by geological and rock mechanics methods. These movements cause horizontal stresses overcoming with 50-60 percents the vertical geostatic stresses.

Extraction technology description

The extraction technology has been worked out by VNIMI "GALYRGIA" - Saint Peterbutg, (Russia). The salt is extracted from above 1200m level by solution using a telescopic borehole system circulating water at a well head pressure of 50 bars. The roof shape is controlled by floating oil layer. The boreholes are situated in square net 200x200m. The shape of the chambers is controlled twice annually. Their diameter varies from 100 to 140m (the design diameter is 50m).

Explotation regime parameters

The observations for the period from 1952 to 1992 show that the salt massif has become lighter with $P_q = 5.3 \cdot 10^6 \text{t}$ dry material - approximately 1500000t per year - fig.4. The relationship between extracted salt mass and the subsidence

Stress tensor variation
 in the salt diapir at depths to 1000m - fig.3a, 1200-1700m -
 fig.3b and underlying formation - fig.3c

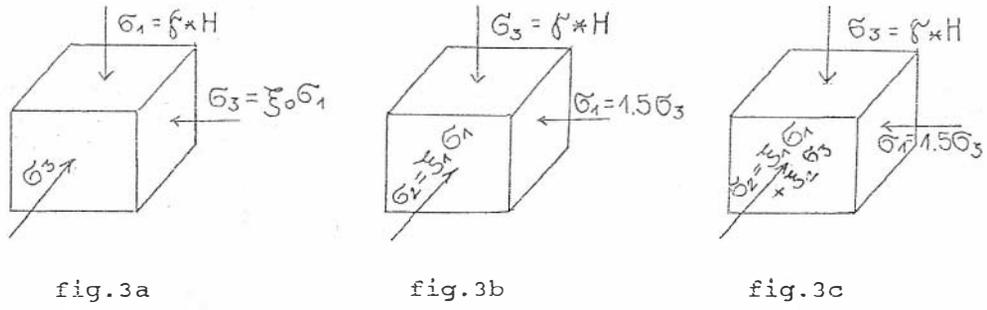


Figure 3

Relationship between extracted salt mass, time,
 subsidence and earthquake occurrence and intensity

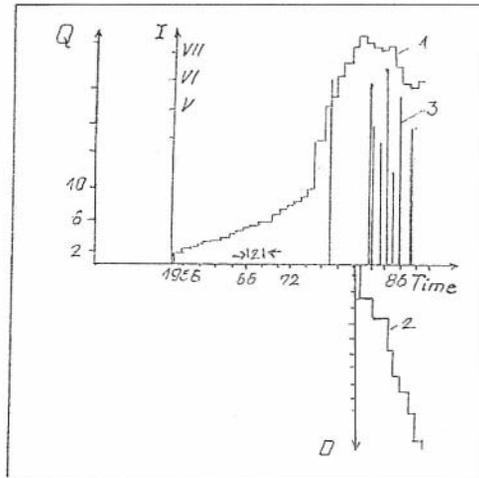


Figure 4

Q - quantity extracted salt mass in 10×10^5 t (1),
 D - subsidence of the area over chamber 6 in mm (2),
 I - earthquake intensity MSK64 (3).

of the area over the chambers is shown on the same figure. The volume of the extracted salt quantity is $23.7 \cdot 10^7$ m.cub. The extraction coefficient of the deposit as a whole is $K=0.1$. The brine weight in all chambers is $P_r=31.88 \cdot 10^6$ t. The total unloading $P_t=P_o-P_r$, directly influencing the stress and strain redistribution in the massif and surrounding rocks and this way influencing the regional seismic activity is approximately $P_t=30 \cdot 10^6$ t. P_o is the chamber material weight before extraction.

Historical seismicity of the region

The Bulgarian territory is characterized by high seismic activity (the intensity map shows $I=VII$ by MSK 64 for more than 65% of the territory). There are many differently orientated faults in the deposit region, forming complex tectonical points. On the potential seismic activity map is shown that near the salt deposit could occur earthquakes with magnitude $M=5.6-6$ at distance $R=10$ km; $M=5.1-5.5$ & $r=1$ km and $M=7.5-8$ & $R=55$ km. Synthetic accelerograms from this potential sources are generated - fig.5. Static and dynamic analysis is carried out using the FEM (finite element method) and these generated accelerograms.

Observation systems results

Two local observing systems are built in the deposit region - geodetical and seismic one. The geodetical observations have been processed on large and close net. A constant subsidence with average velocity 3 - 4.5mm per year and horizontal block movements have been observed in this region for the recent few years.

Seismic regime parameters

More than 40 components have been registered and analyzed for the period of 10 years after the local seismic network building. The number of events $N(M)$ with magnitude M in the range of $1.8 \leq M \leq 4.9$, corrected for one year is $\lg N=1.90-0.68M$. The epicentral distances vary from 6 to 27km. According to the potential seismic sources map the maximum expected magnitude is 5.6-6 for the area of radius $r=10$ km around the site and reduced depth 5-10km. The main conclusion is that these records are saturated with high frequency vibrations. The maximal periods are in the range of 0.085-0.2sec for the vertical components and 0.1- 0.57sec for the horizontal ones. Another special feature is the impacting events character and the low

Generated synthetic accelerograms
Figure 5a and 5b

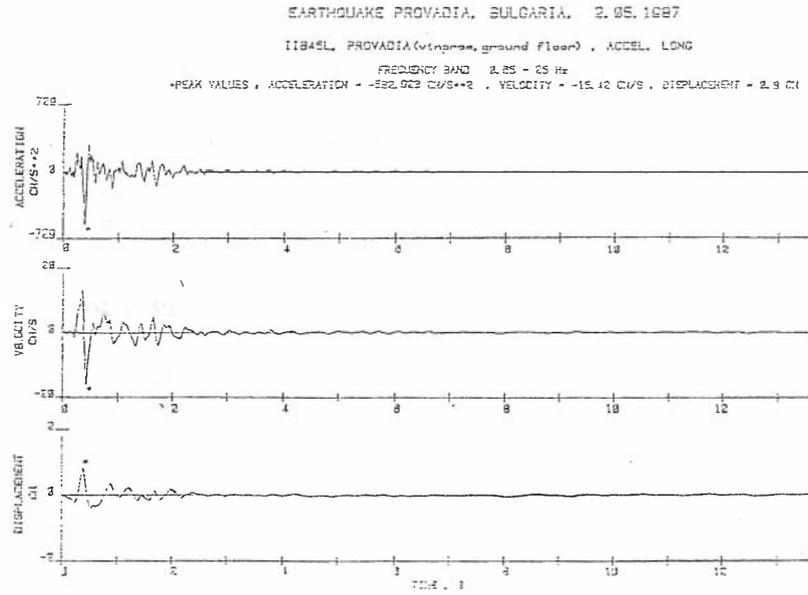
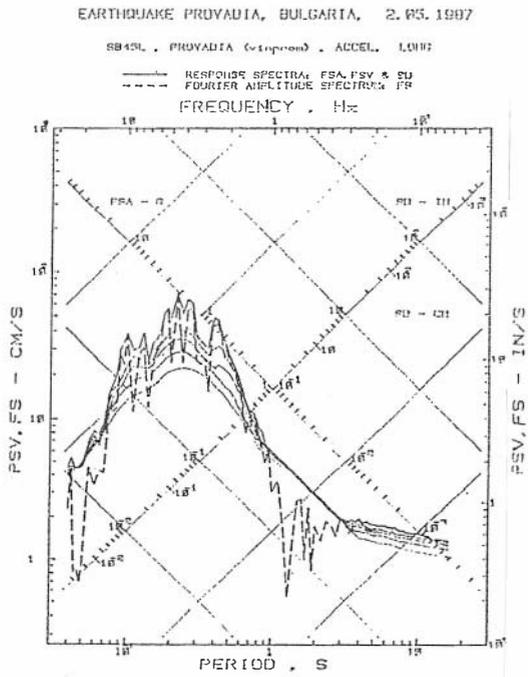


Figure 5a



Beniof graph

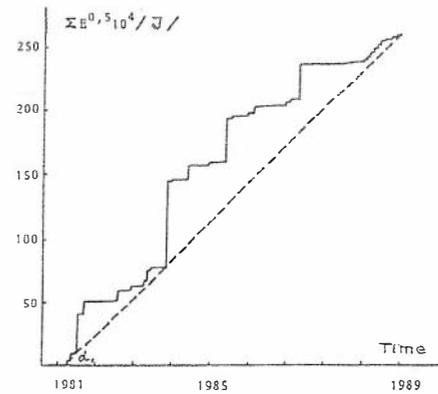


Figure 6

duration of the maximum acceleration part of the records 0.12-2.97sec. The values of the peak ground accelerations are quite high, they overcome 0.5g. The elastic strain release during earthquakes or ground deformations provoking such events can be characterized by Beniof method - fig.6. Energy, accumulated in the terrestrial medium can be expressed by the equation $E=0.5*v*e^2*V$, where v is the Poisson ratio, e is elastic strain and V is terrestrial medium volume, which elastic release causes earthquake. The strain release velocity is characterized by the gradient of the Beniof graph. Analyzing this graph we should have in mind that more accumulated strain energy corresponds to greater limit strength. In our case the gradient of the Beniof graph is 51deg which means that the limit massif strength is relatively low in the discussed area.

Conclusions

Analysis of the stress state field at different depths is processed. It is seen that new conditions of the geodynamical processes, seismic excitations and ground deformations have been created during the process of continuous exploitation of the salt deposit. At large depths, where strong ground motion energy is accumulated, the massif is in a state of unstable equilibrium. Salt extraction could be act as a starting device for energy release. On the other hand the analysis of the observed relationship between seismic activity and extracted salt mass from depths of 1000m shows that an increment of 200000t of the extraction mass can be a starting device for an earthquake occurrence in the case of total extracted mass overcoming 2200000t. The adduced quantitative decisions in our studies give us the reason to suggest that at depths above 1000m less quantity extracted salt could provoke earthquakes. We have to mention that the extracting technology changes could be developed in two main directions - 1) geometrical changes of pillars and chambers sizes and 2) reducing of the well head pressure, but not every change could be useful and quite efficient to reduce seismic activity of the region, having in mind the present region state. That is why we need to specify and improve all deposit observation systems and that way to realize a complex and objective approach for any technological or structural change establishment.

References

1. Konstantinov B., Natural Stress State Analysis of The Salt Rock Deposit in Mirovo, University Edition of the University of Mining and Geology, Department of Engineering Geology and Hydrogeology, Sofia, 1989, pp-128, 141.
2. Houke, Estimating Mohr-Coulomb Friction and Cohesion Values from the Hoek - Brown Failure Criterion, Int.J.Rock. Mech.Min. Sci.&Geomech.Abstr.Vol.27, No3, pp227- 229, 1990.
3. Researching Reports Concerning Geomechanical Conditions of Mirovo Salt Deposit, CLSMEE-BAS, 1987-1991.
4. Knoll P., Fundamentals of a Practical Classification of Mining Induced Seismicity (Rock Bursts), 5th Int. Conf. on Soil Dynamics and Earthquake Eng., Sept.1991, Karlsruhe.
5. Bieniawski Z.T., Geomechanical Classification of Rock Masses and its Application in Tunelling, Proc. 3rd Congr.Int. Soc.Mech.Denver, Vol.2, Part A, pp.27-32.
6. Manev G., Paskaleva I., Kouteva M., The Induced Seismicity Problem in the Mirovo Salt Deposit in Provadia Region, Journal of Mining and Geology, No4, 1992, Sofia.