MECHANISM OF RIGID OVERLAYING OF CARBONIFEROUS STRATA FAILURE IN FACE MINING IN THE CASE OF A MULTISEAMS DEPOSIT

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ABSTRACT

The technique of assessing the failure of rigid overlying strata is applicable to deep mining with thick coal seams (a thickness of more than 1 m in the Ostrava–Karvina Coalfield), and in one case to longwall mining with controlled caving. The assessment of failure of rigid overlying strata makes it possible to distinguish whether the rigid overlying strata of rocks has been deformed or whether a strutting arch has been formed over the goaf below which is an area that is free from stress. Good knowledge of the mining, technical and geological conditions of a given site is a prerequisite for successful evaluation.

There are advantages to utilising surface measurements for interpreting the effects of changes in rock mass, especially in areas of high overlying strata. The practical importance of failure assessment of overlying strata consists in determining the size of the mined-out area where the deformation of the rigid overlying strata occurred, which is dependent on the character of the rock mass.

This paper is set in the context of the expected width of the goaf during deformation of rigid overlying strata within parameters that describe the mining and geological conditions of the locality. Changes in the area of the goaf, based the results of tensometric measurements, will also be placed in context.

KEYWORDS: subsidence trough, rock mass, overlying strata, longwall mining, tensor stress measurement

1. INTRODUCTION

The results of surface subsidence measurements can be used for assessing the failure of the strata overlying the exploited seams. In many cases, a strutting arch is created over the mined out space, and there is no force strong enough to cause failure of the entire thickness of the relatively consistent and rigid overlying strata. When the strutting arch is created, a very large and concentrated rock load can occur, and anomalous geomechanical events may arise. However, even in cases where the intact overlying strata are breaking, the range of the breaking must be precented from extending due to subsequent exploitation. Overhangs of the unfaulted firm layers which are tailed into the non-undermined overlying strata take a part in a considerable surcharge of the affected area. That is why in this area a high stress concentration occurs.

2. LOCALITY DESCRIPTION

This article presents the results of an assessment of the failure of overlying strata from the mining panel 340 206 at the 40th seam in the second block of the Darkov mining area. Mining operations started in July 2011. The extraction is being performed by advanced longwall mining with controlled caving, with an average extracted thickness of 5 m. The advancement of the longwall face and previous mining in the area surrounding the assessed mining panel is presented in Figure 1. The average daily progress of the mining is approximately 3 m. The exploited section of the 40^{th} seam in the second block of the Darkov mining area is located at a depth of 800 m. Previous mining of the 39^{th} seam was undertaken at a depth of 770 m. These seams form part of the Saddle members of the Karviná formation. The general dip of layers in the second block is 5 to 10° to the NNE.

Information about the conditions of the overlying strata was obtained from borehole M4–61. This contains a 26 m thick interlayer between coal seams 39 and 40. The interlayer is mostly comprised of coarse-grained sandstones and conglomerates. The rigidity of the effective overlying strata of the 40th seam is characterized by an inflexibility coefficient value of 10.9. The effective overlying strata of the 40th seam is approximately 30 m thick (six times the extracted thickness). The interlayer between coal seams 37f and 39 is 44 m thick and it consists mainly of coarse and fine-grained sandstone.

The tectonic evolution of the area is very complicated. West of mining panel $340\,206$ there is found the tectonic fault Gabriela, inclined 70° to the



Fig. 1 Position of borehole M4–61 and surface point d13 in relation to mining panel 340 206 and previous mining.



Fig. 2 Contour map of surface subsidence and mined out area of the 340 206 mining panel before October 13th 2011.



Fig. 3 Contour map of surface subsidence and mined out area of the 340 206 mining panel before March 5th 2012.

east and with a drop size of 80 m. East of the planned end of mining panel 340 206 is the tectonic fault Eliška, inclined 70° to the east and with a drop size of 50 m. North of mining panel 340 206 is the tectonic fault Ležatá, inclined 60° to the North and reduced in the size to 30 m.

Only the vertical component of the surface point motion, i.e. the subsidence, is taken into account in modelling the creation of the subsidence depression. The size of the surface point subsidence is determined as the difference between two elevations obtained from periodic surveying. If a network of surface points is used for surface height surveying, it is possible to model the creation of the subsidence depression.

In order to measure movements on the surface, points were fixed in a surface network the position and elevation of which is periodically surveyed by the Satellite (GNSS: Global Navigation System Doležalová et al., 2012; Kajzar et al., 2012). In this case, the geostatistical kriging interpolation method with a linear variogram model was used as a subsidence modelling method. The contour maps in Figures 2 and 3 show the progress of the subsidence depression until October 13th 2011 and March 5th 2012, respectively. A contour map is a 2D representation of 3D data. The first two dimensions are the x and y coordinates and the third dimension is represented by isolines. The relative spacing of the contour lines indicates the relative slope of the surface subsidence.

The gate on the north side of mining panel 340 206 was selected for monitoring the manifestations of longwall face advance at the adjacent mine gates. This gate was equipped with a series of three-level extensometers in the roof, and also with dynamometers, which monitor the process of loading selected anchor reinforcement elements. A special cone gauge probe was installed in the overburden of the selected gate at a chainage of 340 m to monitor stress-induced changes (see Figure 1). The mutual distance of both the perpendicular projections of the gauge probe and point d13 location on the gate directional line was 210 m. An upward borehole with a length of 19 m and diameter of 76 mm, with a special conical-shaped bottom, was used. Consequently, an oriented strain gauge measuring probe was glued into this modified borehole. Stress was detected using the CCBM probe (Compact Conicalended Borehole Monitoring), which is based on accurately measuring the deformations of the conical bottom of the borehole. Then, using laboratory tests which had determined the mechanical properties of materials from the probe's place of installation, these deformations could be transformed into the shape of the stress tensor (Sugawara and Obara, 1999). The values were collected at a frequency of about one week and were compared to the current position of the



Fig. 4 Subsidence of surface point d13, with graphical display of registered significant SL events and stress changes in the vicinity of the goaf.

longwall face. Based on these measurements, dependencies of stress state changes were continuously constructed, i.e. the full shape was calculated of the tensor of the stress changes that were induced by the progress of the longwall face (similar to Staš et al., 2011). For the purposes of this article, the dependence of the change in the selected vertical stress component (negative values represent compression) in relation to the time axis is shown at the bottom of Figure 4.

3. ASSESSMENT OF THE FAILURE OF OVERLYING STRATA

It is necessary to consider the rigid overlaying strata as a stratified nonhomogeneous beam composed of layers that vary in thickness and firmness and which are faulted by a complicated system of tectonic faults. The uneven distribution of the mechanical properties of the overlaying rocks in the stratified mountains also results in their diverse deformations. The overburden is formed of rigid layers with a large bearing capacity, but with only a small capability of deflection. These layers are deformed in a brittle manner after exceeding the strength limit. The overburden is further formed of layers flexibly adapting to changing storage conditions and is capable of great deflection, and also of layers that can adapt to changed conditions in a plastic manner. This means that at the time of the rigid overlaying strata breakthrough, only those layers with a small deflection capability are deformed in a brittle manner. Deformation of flexible layers when a complete failure occurs is not fragile. A fragile failure of these layers occurs only subsequently as part of the mining progress (Jiránková, 2010; Jiránková, 2012).

Records of registered significant seismological (SL) events in the overburden are used when evaluating the failure of overlying strata. Seismic activity is continuously monitored in the Ostrava–Karvina region (OKR). Currently, seismic stations belonging to local and regional networks operate in the area. The local network is formed of stations managed by individual mines to monitor their own mining areas in the Karvina area of OKR. The regional network is formed of stations belonging to the Seismic Polygon of the Green Gas DPB, a.s. company. This Seismic Polygon contains ten stations. Data assessments by the local networks, together with



Fig. 5 Graphical display of significant SL events and advancing coalface before October 13th 2011.

those of the stations of the Seismic Polygon, have been processed since 2002 in the assessment centre at Green Gas DPB, a.s. in Paskov (Holečko, 2008).

Point d13 was selected to present a joint of both surface subsidence assessment and seismological activity during exploitation of mining panel 340 206 (Figure 4). At the bottom of Figure 4 are graphically depicted measurements made of stress changes in the vicinity of the goaf. Registered SL events and the advance of the working face front until October 13th 2011 and until March 5th 2012 are graphically depicted in Figures 5 and 6. The mutual distance of the projection of point d13 and the projection of the stress measurement point, which are both perpendicular to the directional line of gate 340 226, is about 210 m.

Significant seismological events (with an energy greater than 10^4 J) are differentiated by whether they are of a natural character or whether they originate from destress rock blasting (DRB). The first significant SL event was registered 34 days after mining of the 340 206 panel started. On that occasion, the working face front moved forward approximately 60 m in the direction of advance and the width of the working face was approximately 200 m. Another SL event was registered significant SL events were natural and occurred North of the goaf. Then, a series of 11 events, both natural and caused by DRB, occurred between August 11^{th} 2011 and September 22^{nd} 2011. During this period, no significant stress changes were

observed around the goaf (see Figure 4). At the same time, there was a gradual increase in subsidence of the surface points, and there was therefore a gradual failure of the overlying rocks. Natural SL events showed a brittle manner of deformation in some of the overlying strata. However, elastic deformation of the layers, which were flexibly adapting to the changing bedding conditions, were recorded in the layered overlying strata.

On October 10th 2011, DRB was undertaken with an extremely high seismic efficiency, which resulted in rockburst; the mine was so damaged that advance of the working face front had to cease for about five weeks. Subsequently, this resulted in a reduced increase in surface subsidence. During this break in mining, one significant SL event was registered (on October 26th 2011) in the area above the initial cross drift (see Figure 6). The stress change results (measured according to the methodology described in Staš et al., 2007, 2008, 2011, Knejzlík et al., 2008) made it apparent that the vertical component of the stress in the location of the installed probe reacted to DRB with an increase of the compression load, but barely changed when the longwall face advance stopped (Figure 4). A further gradual increase of the compression load in the vicinity of the goaf occurred only with the resumption of mining. The continuation of mining produced a new increase in surface subsidence and was accompanied by significant SL events that originated only from DRB. It can be assumed that at this time elastic



Fig. 6 Graphical display of significant SL events and advancing coalface before March 5th 2012.

deformations of the overlying strata were more dominant than brittle deformations.

The challenge is how to use the experiences of past evaluations from similar localities when the rock mass is highly inhomogeneous and the natural conditions are unique, not only in each area but also in each exploited mining panel. Therefore the assessment of failure of the overlying strata must be primarily based on measured data in the given area. Only then can the result of the assessment be of practical significance.

4. CONCLUSION

It is necessary to consider the rigid overlaying strata as a stratified non-homogeneous beam composed of layers varying in thickness and geomechanical conditions. The method of assessing failure of the strata failure overlying extracted coal seams is based upon the simultaneous assessment of surface subsidence and seismological activity, having regard to the spatio-temporal progress of mining and to the character of the rock mass. This means that at the time of breakthrough of the rigid overlaying strata, only those layers that have a small capability of deflection undergo brittle deformation. In contrast, the layers that have are better able to deflect flexibly can adapt to the changing bedding conditions at the time of failure of the rigid overlaying strata.

The development of the surface subsidence presented by the subsidence curve of point d13, together with records of registered significant seismological events and previous experience (Jiránková et al., 2012), show that there was a gradual deformation of overlying strata during the extraction of mining panel 340 206. The apparent onset of surface subsidence is clear in Figure 2, accompanied by the occurrence of both natural seismological events and seismological events related to DRB (destress rock blasting). At this time (from August to October 2011), a gradual deformation of the rigid overlying strata occurred. After DRB was performed on October 10th 2011, the advance of the working face front ceased for about five weeks, which resulted in a reduced increase in surface subsidence. Additionally, it is apparent from the stress change measurement results that the vertical component of the stress in the location of the installed probe reacted to DRB with an increase of the compression load, but it hardly changed when the longwall face advance stopped.

The following mining activity brought a new increase in surface subsidence and was accompanied by significant seismological events that originated only from DRB. It can be assumed that at this time elastic deformations of the overlying strata dominated brittle deformations.

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