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SINGULAR SPECTRUM ANALYSIS FOR HETEROGENEOUS TIME SERIES BY TAKING ITS FORMAL ERRORS INTO ACCOUNT

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ABSTRACT

Since the formal errors of many geodetic time series are also available, this paper proposes a new singular spectrum analysis (SSA) approach with stabilizing weights (SSAW) by taking the formal errors into account, where the weight of time series data is constructed based on the ratio of formal error to signal power spectrum. The formulae of the proposed SSAW are derived in detail and then used to process the real Global Mean Sea Level (GMSL) time series compared to the traditional SSA. When the first 10 principal components are used to fit the GMSL time series, the fitting errors of the SSAW and traditional SSA are 4.80 mm and 5.14 mm, with the reduction of 6.61 %. According to the 500 simulations based on the reconstructed signals and formal errors of GMSL time series, the mean root mean squared errors and mean absolute errors of reconstructed signals using the SSAW relative to traditional SSA are reduced from 2.18 mm to 1.66 mm and 1.67 mm to 1.34 mm, respectively. Therefore, if the formal errors of a noisy time series given, the proposed SSAW approach is suggested to analyze this time series.

1. INTRODUCTION

Singular Spectrum Analysis (SSA) has been successfully applied to extract the trend and periodic signals from a time series without prior knowledge (Broomhead and King, 1986; Vautard and Ghil, 1989; Golyandina et al., 2001). In SSA a trajectory matrix is first constructed from a time series, and then principal component analysis is carried out to pick out dominant components. The signals of the time series are reconstructed with the dominant components and the remaining components are regarded as the noise (Sivapragasam et al., 2001; Vitanov et al., 2008). SSA has been widely used in geosciences, e.g. the stream flow and sea-surface temperature analysis (Robertson and Mechoso, 1998; Kondrashov and Ghil, 2006; Chen and Sneeuw, 2016; Zotov et al., 2017), the seismic tomography (Oropeza and Sacchi, 2011; Tiwari and Rajesh, 2014; Rajesh and Tiwari, 2014), earth rotation and climate variability (Zotov et al., 2016), the gravity anomaly separation (Kumar et al., 2017) and monthly gravity field analysis (Zotov and Shum, 2010; Rangelova et al., 2012). For processing an incomplete time series, several SSA approaches, such as spatiotemporal data filling (Kondrashov and Ghil, 2006), iterative data filling (Kondrashov et al., 2010), or compensation via a scale (Schoellhamer, 2001) and its improved version (Shen et al., 2015), which is based on the principle that a time series can be reconstructed with its principal components (Shen et al., 2014).

SSA can be directly used to process the time series contaminated by additive noise (Allen and Smith, 1997; Wu and Chau, 2011). When a time series is contaminated by multiplicative noise, the traditional strategy first converts the multiplicative noise to additive noise by taking log-transform (Golyandina et al., 2001), then processes it with traditional SSA. SSA outperforms other methods in extracting signals from heterogeneous time series, such as the Fourier transform, wavelet transform, empirical mode decomposition and Kalman filters (Kumar et al., 2017; Barrios-Muriel et al., 2016; Watson, 2016; Chen et al., 2016). The covariance or formal errors of a time series are usually not being considered in most SSA approaches, however, formal errors are available for many geodetic time series data. Some SSA approaches have been developed by prewhitening a time series data using its noise matrix (Cadzow, 1988) or rank-deficient covariance matrix (Hansen et al., 2005). Nevertheless, the prewhitening processing will change the structure of original time series. In order to keep the signal structure unchanged the weighted singular value decomposition (Gillard and Zhigljavsky, 2016) and weighted hankelization (Zvonarev and Golyandina, 2017) applied the weights to the trajectory matrix rows and columns, which are thereby named as weighted SSA approach. However since signals are unknown, an iterative algorithm is needed. To avoid iteration, a new SSA approach with stabilizing weights (SSAW) is developed in this

contribution, in which the weight is constructed based on the ratio of formal error to signal power spectrum.

The rest of this paper is organized as follows: the SSAW approach is introduced after traditional SSA briefly presented in Section 2. The sea level change signals are extracted from Global Mean Sea Level (GMSL) time series and analyzed in Section 3. Then, synthetic examples are carried out to demonstrate the performance of the SSAW compared to traditional SSA in extracting signals in Section 4, and conclusions are given in Section 5.

2. METHODOLOGY

2.1. SINGULAR SPECTRUM ANALYSIS

From a time series $x_i (1 \leq i \leq N)$, we can construct an $L \times K$ ($K=N-L+1$) trajectory matrix X with a window size L (Vautard et al., 1992) as follows,

$$X = \begin{bmatrix} x_1 & x_2 & \cdots & x_{i+1} & \cdots & x_K \\ x_2 & x_3 & \cdots & x_{i+2} & \cdots & x_{K+1} \\ \vdots & \vdots & \vdots & \vdots & \vdots & \vdots \\ x_L & x_{L+1} & \cdots & x_{i+L} & \cdots & x_N \end{bmatrix} \quad (1)$$

Its lagged covariance matrix C is formulated as $C = XX^T$. Then the matrix C is decomposed as,

$$C = V\Lambda V^T \quad (2)$$

where Λ is a diagonal matrix with eigenvalues sorted in descending order, its k th number is denoted as $\lambda_k (1 \leq k \leq L)$. V is an orthogonal matrix; v_k is its k th eigenvector. The principal component matrix A is determined by,

$$A = VX \quad (3)$$

The k th row vector a_k of A is called the k th Principle Components (PCs). Then the i th element $a_{k,i}$ of a_k is computed with,

$$a_{k,i} = \sum_{j=1}^L x_{i+j-1} v_{k,j}, \quad 1 \leq i \leq N-L+1 \quad (4)$$

where $v_{k,j}$ is the j th element of v_k . Then the Reconstructed Components (RCs) of the time series with the k th PC are represented as (Vautard et al., 1992),

$$x_i^k = \begin{cases} \frac{1}{i} \sum_{j=1}^i a_{k,i-j+1} v_{k,j} & 1 \leq i \leq L-1 \\ \frac{1}{L} \sum_{j=1}^L a_{k,i-j+1} v_{k,j} & L \leq i \leq N-L+1 \\ \frac{1}{N-i+1} \sum_{j=i-N+L}^L a_{k,i-j+1} v_{k,j} & N-L+2 \leq i \leq N \end{cases} \quad (5)$$

where x_i^k is the i th element reconstructed from k th PC. Since the eigenvalues of covariance matrix C are sorted in descending order, the signal of the time

series can be represented by the first several RCs, while the remaining RCs being treated as noise. Thus the signals of the time series are the sum of dominant RCs as follows,

$$\hat{x}_i = \sum_{k=1}^n x_i^k, \quad i=1,2,\dots,N \quad (6)$$

where \hat{x}_i denotes the i th element of reconstructed signals from the first n dominant RCs.

2.2. SINGULAR SPECTRUM ANALYSIS WITH STABILIZING WEIGHTS

In traditional SSA approach the Equations (1)~(6) are normally used to process a time series without considering its covariance or formal errors, which implies that the priori information of observational noise is neglected in traditional SSA approach. In order to use the priori covariance Q , Hansen et al (2005) proposed one solution generating a new time series by prewhitening the original time series as,

$$x' = xQ^{-1/2} \quad (7)$$

where $x = (x_1, x_2, \dots, x_n)$ is a time series represented in row vector form. It is obvious that the covariance of the generated time series x' is an identity matrix. Therefore, Equations (1)~(6) can be applied to extract the signals \hat{x}_i from x' . Then the signals of original time series are recovered as,

$$\hat{x} = \hat{x}'Q^{1/2} \quad (8)$$

Since not the full covariance Q but only its diagonal elements are available for most geodetic time series, the Equations (7) and (8) can be simplified as,

$$x'_i = x_i \frac{\sigma_0}{\sigma_i} \quad (9)$$

and

$$\hat{x}_i = \hat{x}'_i \frac{\sigma_i}{\sigma_0} \quad (10)$$

where the formal error σ_i is the squared root of the i th diagonal element of Q , σ_0 is the formal error of unit weight which is computed with the principle of keeping the total energy of the time series unchanged (Li et al., 2015). However, the prewhitening processing in Equation (7) or (9) seriously changes the signal structure of the original time series, especially when σ_i is varied significantly from epoch to epoch, since the transformation of Equation (9) implies that both the signals and noises of the time series are multiplied by a weight factor σ_0 / σ_i . In the weighted SSA, an iterative procedure is needed (Gillard and Zhigljavsky, 2016; Zvonarev and Golyandina, 2017). For this reason, a new weight factor is derived in the following parts, which can suppress noise, but hold more signals than Hansen et al. (2005).

A time series $x_i (1 \leq i \leq N)$ is composed of signal s_i and noise e_i as follows,

$$x_i = s_i + e_i \quad (11)$$

where s_i is stationary with mostly low frequencies in the spectrum, e_i consists of independent samples with slowly changing variances.

Supposing the signal and noise are independent, the variance of x_i can be derived via the law of error propagation law as,

$$\sigma_{x_i}^2 = \sigma_{s_i}^2 + \sigma_{e_i}^2 \quad (12)$$

where $\sigma_{s_i}^2$ is the variance of signal s_i . If both the σ_{e_i} and σ_{s_i} are available, we can convert the original time series as,

$$x'_i = \sigma_{s_i} x_i / \sqrt{\sigma_{s_i}^2 + \sigma_{e_i}^2} = x_i / \sqrt{1 + \sigma_{e_i}^2 / \sigma_{s_i}^2} = x_i / \alpha_i \quad (13)$$

where the weight factor $\alpha_i = \sqrt{1 + \sigma_{e_i}^2 / \sigma_{s_i}^2}$, $i = 1, 2, \dots, N$ can suppress the impacts of noise. By the law of error propagation, we can derive from Equation (13) as,

$$\sigma_{x'_i}^2 = \sigma_{s_i}^2 \quad (14)$$

Therefore, the converted time series x'_i can have the same variance as the original signals. With the converted time series x'_i or the correspondent trajectory matrix X' , one can form a lag covariance C' as,

$$C' = X' X'^T \quad (15)$$

The eigenvectors $v'_{k,j}$ of the lag covariance C' are then used to calculate the PCs $a'_{k,i}$ with,

$$a'_{k,i} = \sum_{j=1}^L x'_{i+j-1} v'_{k,j}, 1 \leq i \leq N - L + 1 \quad (16)$$

Then the RCs of the time series are calculated as Equation (5) by replacing the $a_{k,i}$ and $v_{k,j}$ with $a'_{k,i}$ and $v'_{k,j}$. With the signals \hat{x}' of the converted time series determined with Equation (6), the signals of the original time series is recovered with,

$$\hat{x}_i = \hat{x}'_i \alpha_i \quad (17)$$

Since a new weight factor $\alpha_i = \sqrt{1 + \sigma_{e_i}^2 / \sigma_{s_i}^2}$, which is based on the ratio of σ_{e_i} to σ_{s_i} , is introduced in our developed approach, we call it as SSAW. The weight factor α_i is larger than 1 and increases with respect to the ratio of σ_{e_i} to σ_{s_i} . Hence the larger the noise is, the bigger the weight factor will be. Therefore, the impacts of heterogeneous noise are

suppressed and the signal to noise ratio will increase in SSAW processing. When the signal variance is not known, it should be estimated beforehand. Especially, when the time series is stationary, the signal variances $\sigma_{s_i}^2$ of all epochs are equal, which can be estimated by using the reconstructed signals,

$$\hat{\sigma}_s^2 = \frac{1}{N-1} \sum_{i=1}^N (s_i - \bar{s})^2 \quad (18)$$

where s_i is the reconstructed signal at i th epoch; \bar{s} is the mean signals of all epochs; N is the number of epochs of the time series.

2.3. WINDOW SIZE AND RECONSTRUCTED ORDER

The window size L and number of dominant PCs are the two key factors to be determined in SSA. The window size mostly depends on the research topic and the priori information of the time series (Golyandina and Zhigljavsky, 2013). L should be large enough theoretically but not larger than $N/2$ (Golyandina, 2010). If a time series contains a periodic component, the window size had better to be proportional to the period. If one is interested in signal as a whole, but not the separability among signal components, the window size is not a key factor (Golyandina et al., 2001). There are different approaches to determine the number of dominant PCs for reconstructing signals, such as the threshold for the sum contribution of dominant PCs (Figueiredo et al., 2011; Ma et al., 2011), the first inflection point in the logarithm of eigenvalues diagrams (Hu et al., 2009), and the weighted correlations (w-correlation) (Golyandina et al., 2001). Due to its intuitive, w-correlation matrix is widely used to detect the signal subspace (Golyandina et al., 2018). Thus the subspace of dominant PCs is determined with w-correlation matrix. The weights for computing weighted correlations are determined by,

$$w_i = \begin{cases} i, & 0 \leq i < L \\ L, & L \leq i \leq K \\ N - i + 1, & K < i \leq N \end{cases} \quad (19)$$

From Equation (19), it is apparent that the weight is equal to the number of the element x_i in the trajectory matrix X . Define the inner product of two components PC^i and PC^j of length N is,

$$(PC^i, PC^j)_w \stackrel{\text{def}}{=} \sum_{k=1}^N w_k PC_k^i PC_k^j \quad (20)$$

The weighted correlation to measure the degree of approximate separability between PC^i and PC^j is defined as follows (Golyandina et al., 2001),

$$\rho_{i,j}^{(w)}(PC^i, PC^j) \stackrel{\text{def}}{=} \frac{(PC^i, PC^j)_w}{\|PC^i\|_w \|PC^j\|_w} \quad (21)$$

If the weighted correlation is 0, it means that the components PC^i and PC^j are fully separable. If the weighted correlation between two components PC^i and PC^j is larger than a certain threshold, then the two components are regarded as non-separable. Therefore, the PCs for reconstructing signals are determined based on the threshold ε as,

$$\max(|\rho_{i,j}^{(w)}(PC^i, PC^j)|) < \varepsilon, \quad i = 1, 2, \dots, L; j = i + 2, \dots, L \quad (22)$$

If the i th component doesn't satisfy the Equation (22), then the reconstructed order is fixed at $i - 1$.

3. GLOBAL MEAN SEA LEVEL TIME SERIES ANALYSIS USING THE SSAW

The GMSL time series presented in Figure 1 includes 1368 monthly averaged data spanning from 1900 to 2013, and the corresponding formal errors are presented in Figure 2, which are available from the web site (<http://www.psmsl.org>). Since the number of tide gauge stations continually increases all over the world, the formal errors significantly become smaller from 1900 to 2005. However, the formal errors become larger from 2005 to 2013, mainly for fewer records are available in these years due to the delay in the transmission of GMSL information to the PSMSL (Church and White, 2011). For GMSL time series, the trend, which contains more than 90 % of the total variance, is pre-determined using least squares fit based on quadratic polynomial functional model (Fig. 1). The de-trended GMSL time series (after removing the trend) is shown in Figure 3 with blue solid line, which is then processed with traditional and SSAW and finally the trend is added back to the reconstructed signal.

The window size is chosen as $L=600$, which is equal to 50 years, less than half of the length of de-trended GMSL time series through experiments comparison. The w-correlations are employed to determine the PCs for reconstructing de-trended GMSL signals, and the 10 leading PCs are determined when the threshold is chosen as 0.5 empirically. Thereby, 10 leading PCs (same as in SSA) are used to reconstruct signals in the SSAW. With the reconstructed de-trended GMSL signals from 10 leading PCs, we can estimate that $\hat{\sigma}_s^2 = 58.46 \text{ mm}^2$. The converted de-trended GMSL time series by Equation (13) is shown in Figure 3 with red solid line.

The w-correlations between any pair of the 30 leading RCs are depicted in Figures 4 and 5 for traditional and SSAW, respectively. It is expected that the indicated pairs are 1-2, 3-4, 5-6, 7-8 and 9-10, so the RCs 1-10 are as signals and the RCs 11-600 are treated as noise. The noise and its frequency spectrum are presented in Figure 6, in which the noise power spectrum is about 10, much less that of signal.

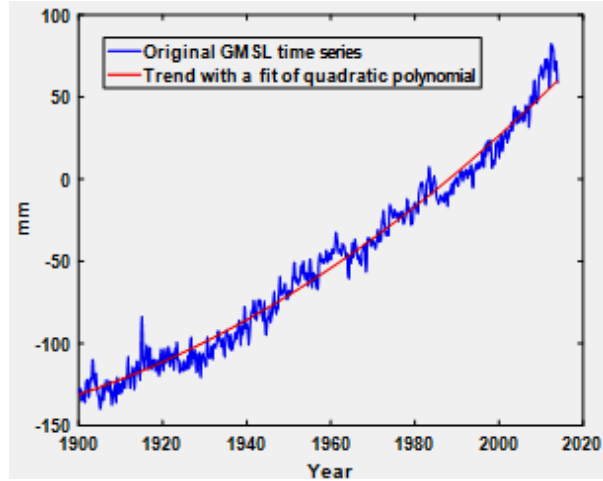


Fig. 1 Original GMSL time series and corresponding trend term.

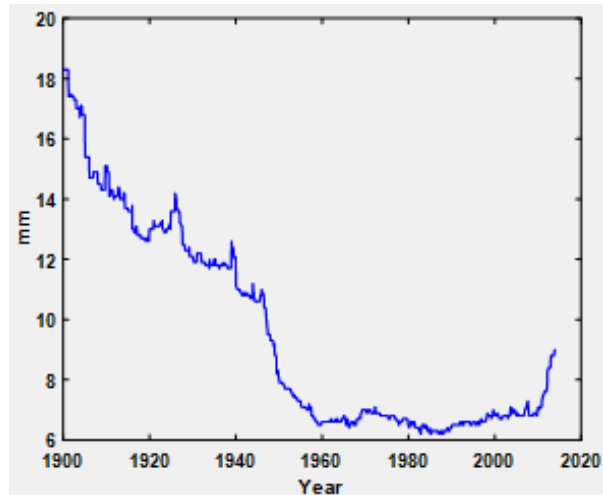


Fig. 2 Formal error series.

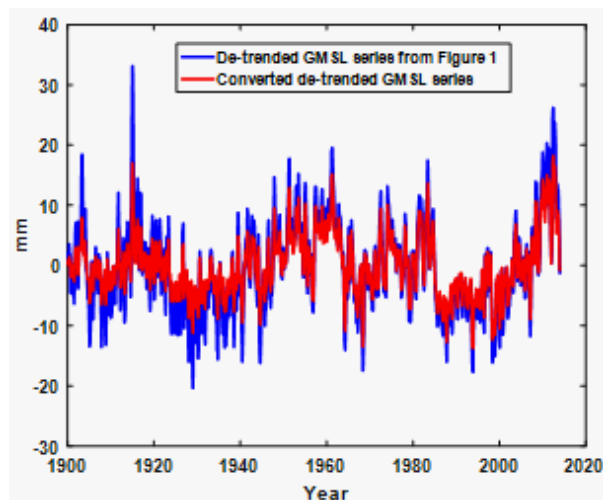


Fig. 3 The de-trended and its converted GMSL time series.

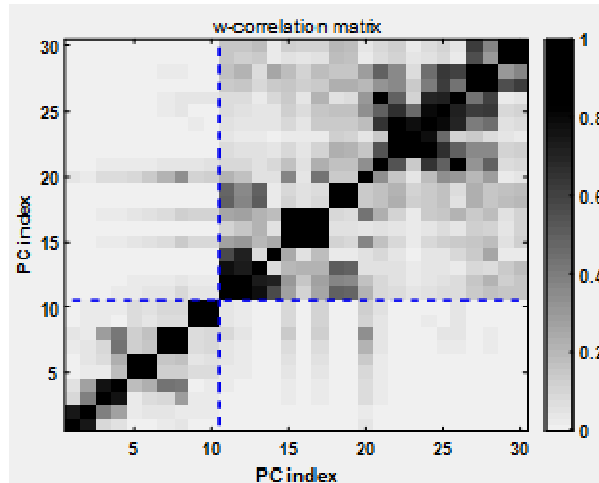


Fig. 4 The w-correlations of 30 leading RCs of traditional SSA.

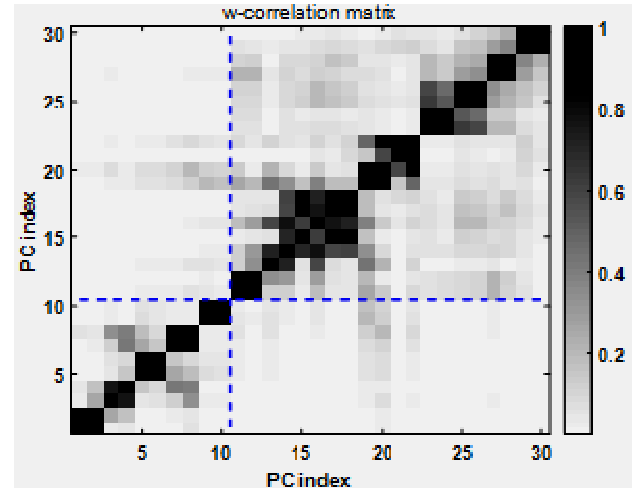


Fig. 5 The w-correlations of 30 leading RCs of SSAW.

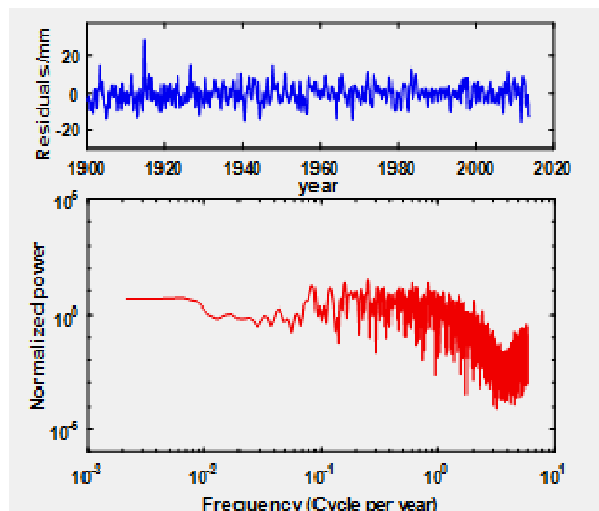


Fig. 6 The noise (RCs 11-600) of SSA (top) and its frequency spectrum (bottom).

The periodograms of the 5 indicated pairs (1-2), (3-4), (5-6), (7-8) and (9-10) are presented in Figure 7, which reveal the impacts of Pacific Decadal Oscillation (PDO), El Niño/South Oscillation (ENSO) and sunspots on the GMSL time series, since the periods of PDO, ENSO and sunspots are respectively about 20~30, 20~25 and 10~11 years (Chen et al., 1998; Nerem, 1995). The periods of the indicated pairs (3-4 and 7-8) are possibly caused by PDO and ENSO, respectively; the period of harmonic pair (5-6) is consistent with the period of sunspots. Besides, a 57-year oscillation (1-2) is significant in de-trended GMSL time series, which is in a good agreement with Church and White (2011) and Chambers et al. (2012). The periods of 5 indicated pairs of the two approaches are listed in Table 1 together with their amplitudes, which are computed with least squares fit using harmonic analysis. The results show that the differences of periods between the two approaches are very small, but differences of amplitudes are very significant.

The percentages of total variance of 10 leading PCs are shown in Figure 8 for the converted and original GMSL time series, where the PCs of the converted time series contain more energy than that of the original GMSL time series. Specially, the variance of 10 leading PCs of the converted time series is

Table 1 Periods and mean amplitudes of paired harmonic components.

PC index	Traditional SSA		SSAW	
	Period/year	Amplitude/mm	Period/year	Amplitude/mm
1-2	57.00	4.89	57.00	5.19
3-4	30.40	2.72	30.40	2.40
5-6	10.36	2.10	10.36	2.29
7-8	19.00	1.59	19.00	1.97
9-10	7.12	1.17	7.24	1.25

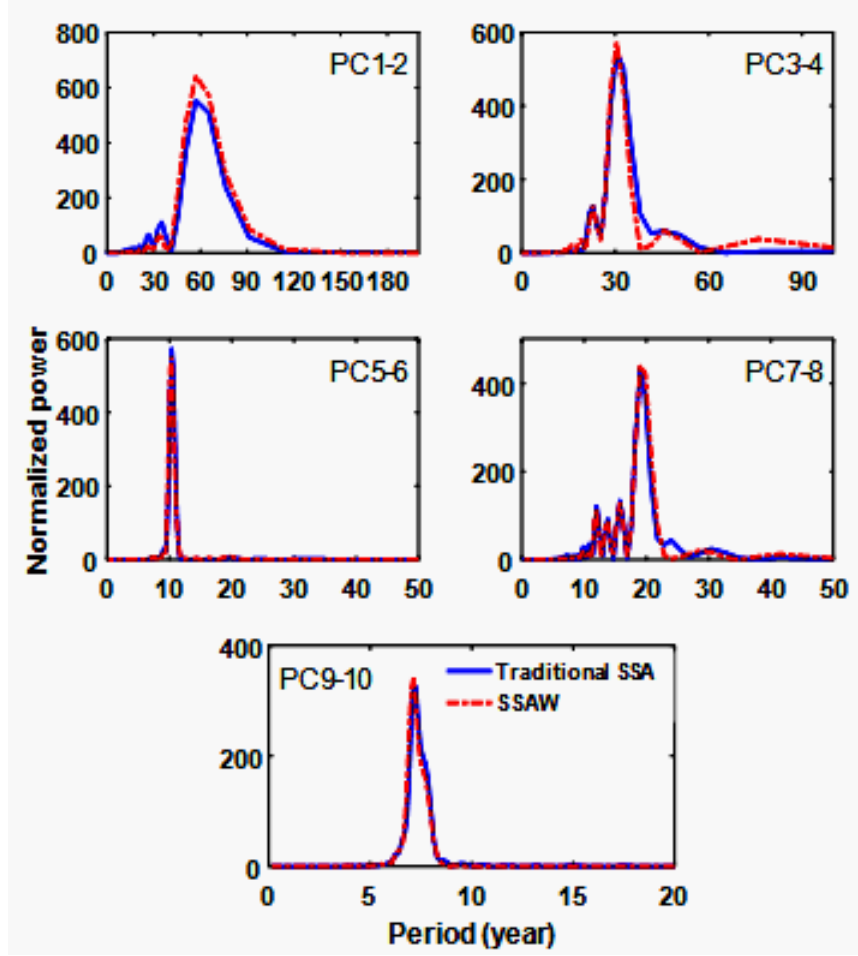


Fig. 7 The corresponding period of the paired harmonic components.

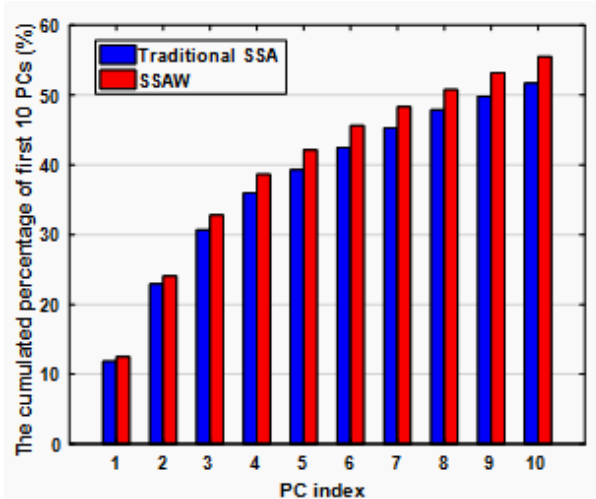


Fig. 8 The cumulated percentage of signal variance for first 10 PCs.

55.5 % of the total variance, much larger than that (51.3 %) of the original time series.

The signals, which are reconstructed from the first 10 PCs with both the SSAW and traditional SSA, are presented in Figure 9 and their differences in bottom panel. It is apparent that the differences are clearly correlated with the formal errors of the time series. If the weights of the SSAW are determined in such a way that the sum of the weights is equal to the epochs number N , the root mean squared error of unit weight indicates the fitting error of the SSAW. According to the definition of weight, the weight of the GMSL time series is defined as,

$$p_i = \sigma_0^2 / \alpha_i^2, \quad i = 1, 2, \dots, N \quad (23)$$

where σ_0^2 is the variance of unit weight, σ_i is the formal error at i th epoch. Since $\sum_{i=1}^N p_i = N$, we can solve σ_0^2 from Equation (23), that is,

$$\sigma_0^2 = N / \sum_{i=1}^N (1 / \alpha_i^2) \quad (24)$$

Substituting (24) into (23) yields the weight as,

$$p_i = N / (\alpha_i^2 \sum_{k=1}^N (1 / \alpha_k^2)) \quad (25)$$

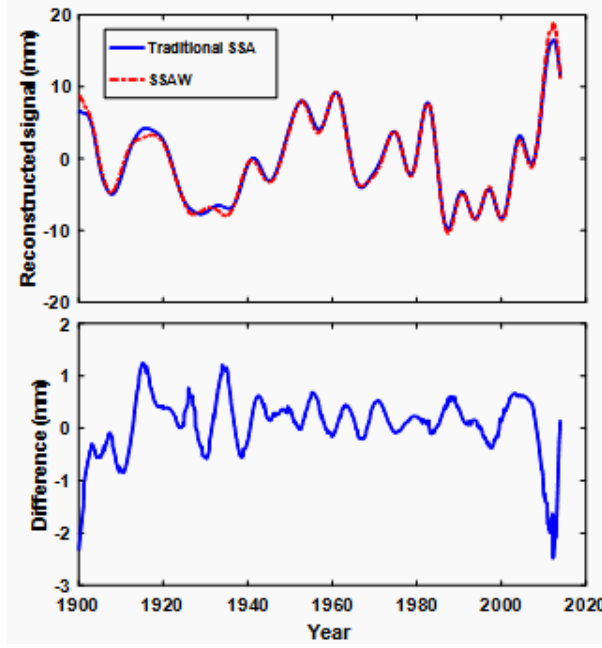


Fig. 9 Extracted signals of the traditional and SSAW (top) and their difference (bottom).

If the signals reconstructed with the SSAW at i th epoch is expressed with \hat{x}_i^w , the fitting error of the SSAW is estimated by,

$$m_w = \sqrt{\frac{1}{N} \sum_{i=1}^N p_i (x_i - \hat{x}_i^w)^2} \quad (26)$$

Similarly for traditional SSA, the unit weight adopted, the fitting error is estimated by,

$$m_s = \sqrt{\frac{1}{N} \sum_{i=1}^N (x_i - \hat{x}_i^s)^2} \quad (27)$$

where \hat{x}_i^s is the reconstructed signal with traditional SSA at i th epoch.

The fitting errors of the GMSL time series using 10 PCs are 4.80 mm and 5.14 mm for the SSAW and traditional SSA, respectively. Compared to traditional SSA, the SSAW can reduce the fitting error about 6.61 %. Therefore, when the formal errors of a noisy time series are available, the SSAW is suggested to deal with the time series.

4. SIMULATION STUDIES BASED ON THE RECONSTRUCTED GMSL SIGNALS

To demonstrate the performance of the SSAW, we carry out simulation studies based on the real signals shown in Figure 9 extracted by traditional SSA approach. The synthetic time series are generated for 500 times by Equation (11) where noise is the normally distributed noise with zero mean and variance σ_i^2 . The real σ_i^2 are the formal errors of GMSL time series shown in Figure 2. The same

window size ($L = 600$) is chosen and first 10 PCs are selected to reconstruct signals. Since the real signal is known, we can compute the root mean squared error (RMSE) and mean absolute error (MAE) with the differences between the reconstructed signal and simulated real signal. They are defined as the following equations:

$$\text{RMSE} = \sqrt{\frac{1}{N} \sum_{i=1}^N (s_i - \hat{x}_i)^2} \quad (28)$$

$$\text{MAE} = \frac{1}{N} \sum_{i=1}^N |s_i - \hat{x}_i| \quad (29)$$

where \hat{x}_i denotes the reconstructed signal at i th epoch, N is the number of observations. The relative improvements (IMP) of RMSE and MAE of the SSAW with respect to traditional SSA is computed with,

$$\text{IMP}_{\text{RMSE}} = \frac{\text{RMSE}_s - \text{RMSE}_w}{\text{RMSE}_s} \times 100\% \quad (30)$$

$$\text{IMP}_{\text{MAE}} = \frac{\text{MAE}_s - \text{MAE}_w}{\text{MAE}_s} \times 100\% \quad (31)$$

where the subscript 'w' and 's' represent the values of the SSAW and traditional SSA, respectively.

The RMSEs and MAEs for the SSAW and traditional SSA are presented in Figure 10, where all the RMSEs and MAEs of the SSAW are smaller than those of traditional SSA. The average RMSE and

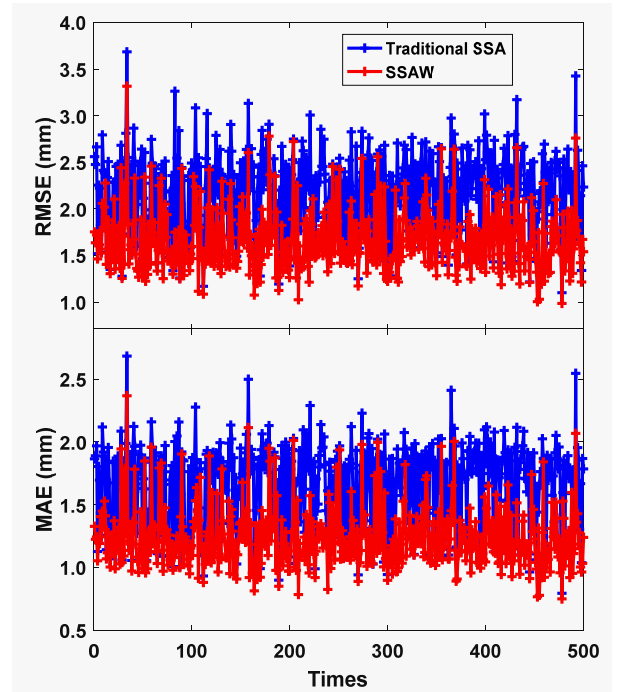


Fig. 10 RMSEs and MAEs of 500 simulations of the SSAW and traditional SSA.

MAE are 1.67 mm and 1.34 mm for the SSAW, 2.18 mm and 1.66 mm for traditional SSA, respectively. The corresponding IMPs of RMSE and MAE are 23.85 % and 19.76 %, respectively. Therefore, the proposed SSAW effectively extracts signals from noisy series compared to traditional SSA.

5. CONCLUSIONS

In order to more efficiently extract signals from noisy time series, the SSAW approach is proposed by considering the formal errors of the time series. The GMSL time series are processed by using the SSAW and traditional SSA; the results demonstrate that the SSAW achieves better performance than traditional SSA in reconstructing signals. The first 10 PCs of the SSAW and traditional SSA contain 55.5 % and 51.7 % of total variance, respectively. When the first 10 PCs are used to reconstruct signals for fitting the original time series, the fitting error of the SSAW is 4.80 mm, much smaller than 5.14 mm, the fitting error of traditional SSA. Moreover, according to the reconstructed GMSL signals and formal errors of the GMSL time series, 500 simulations are carried out. The statistics show that the mean RMSE and MAE of the SSAW relative to traditional SSA are reduced from 2.18 mm to 1.67 mm, 1.66 mm to 1.34 mm, with improvements 23.85 % and 19.76 %, respectively. All the RMSEs and MAEs of the SSAW are smaller than those of traditional SSA. Therefore, when the formal errors of a time series are available, the SSAW approach is suggested to be applied to process this time series.

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REFERENCES

- Allen, M. and Smith, L.: 1997, Optimal filtering in singular spectrum analysis. *Phys. Lett. A*, 234, 6, 419–428. DOI: 10.1016/s0375-9601(97)00559-8
- Barrios-Muriel, J., Romero F., Alonso, F. J. and Gianikellis, K.: 2016, A simple SSA-based de-noising technique to remove ECG interference in EMG signals. *Biomed. Signal Proces. Control*, 30, 117–126. DOI: 10.1016/j.bspc.2016.06.001
- Broomhead, D. and King, G.: 1986, Extracting qualitative dynamics from experimental data. *Physica D: Nonlinear Phenomena*, 20, 2–3, 217–236. DOI: 10.1016/0167-2789(86)90031-X
- Cadzow, J.: 1988, Signal enhancement — a composite property mapping algorithm. *IEEE Trans. Speech Audio Process*, 36, 1, 49–62. DOI: 10.1109/29.1488
- Chambers, D., Merrifield, M. and Nerem, R.: 2012, Is there a 60-year oscillation in global mean sea level? *Geophys. Res. Lett.*, 39, L18607. DOI: 10.1029/2012GL052885
- Chen, J., Wilson, C., Chambers, D., Nerem, R. and Tapley, B.: 1998, Seasonal global water mass budget and mean sea level variations. *Geophys. Res. Lett.*, 25, 19, 3555–3558. DOI: 10.1029/98GL02754
- Chen, Q. and Sneeuw, N.: 2016, Singular spectrum analysis for modeling geodetic time series. *IAG Symposia*. DOI: 10.1007/1345_2016_248
- Church, J. and White, N.: 2011, Sea-level rise from the late 19th to the early 21st century. *Surv. Geophys.*, 32, 4–5, 585–602. DOI: 10.1007/s10712-011-9119-1
- Figueiredo, M., Almeida, A. and Ribeiro, B.: 2011, Wavelet decomposition and singular spectrum analysis for electrical signal denoising. *Proc. IEEE Int. Conf. on Systems, Man and Cybernetics*, 3329–3334. DOI: 10.1109/ICSMC.2011.6084183
- Gillard, J. W. and Zhigljavsky, A.: 2016, Weighted norms in subspace-based methods for time series analysis. *Numer. Linear Algebr. Appl.*, 23, 5, 947–967. DOI: 10.1002/nla.2062
- Golyandina, N.: 2010, On the choice of parameters in singular spectrum analysis and related subspace-based methods. *Stat. Interface*, 3, 3, 259–279.
- Golyandina, N., Korobeynikov, A. and Zhigljavsky, A.: 2018, *Singular spectrum analysis with R*. Springer, 270 pp. DOI: 10.1007/978-3-662-57380-8
- Golyandina, N., Nekrutkin, V. and Zhigljavsky, A.: 2001, *Analysis of time series structure: SSA and related techniques*. Chapman & Hall/CRC, Boca Raton. DOI: 10.1201/9781420035841
- Golyandina, N. and Zhigljavsky, A.: 2013, *Singular spectrum analysis for time series*. Springer, 119 pp. DOI: 10.1007/978-3-642-34913-3
- Hansen P. and Jensen, S.: 2005, Prewhitening for rank-deficient noise in subspace methods for noise reduction. *IEEE Trans. Signal Process*, 53, 10, 3718–3726. DOI: 10.1109/TSP.2005.855110
- Hu, B., Li, Q. and Smith, A.: 2009, Noise reduction of hyper-spectral data using singular spectral analysis. *Int. J. Remote Sens.*, 30, 9, 2277–2296. DOI: 10.1109/TSP.2005.855110
- Kondrashov, D. and Ghil, M.: 2006, Spatiotemporal filling of missing points in geophysical data sets. *Nonlinear Proc. Geoph.*, 13, 2, 151–159. DOI: 10.5194/npg-13-151-2006
- Kondrashov, D., Shprits, Y. and Ghil, M.: 2010, Gap filling of solar wind data by singular spectrum analysis. *Geophys. Res. Lett.*, 37, 15, L15101, 78–82. DOI: 10.1029/2010GL044138
- Kumar, K., Rajesh, R. and Tiwari, R.: 2017, Regional and residual gravity anomaly separation using the singular spectrum analysis-based low pass filtering: a case study from Nagpur, Maharashtra, India. *Explor. Geophys.*, 49, 3, 398–408. DOI: 10.1071/EG16115
- Li, W., Shen, Y. and Li, B.: 2015, Weighted spatiotemporal filtering using principal component analysis for analyzing regional GNSS position time series. *Acta Geod. Geophys.*, 50, 4, 419–436. DOI: 10.1007/s40328-015-0100-1
- Ma, J., Li, Z. and Wang, B.: 2011, Application of singular spectrum analysis to the noise reduction of intrusion detection alarms. *J. Computers*, 6, 8, 1715–1722. DOI: 10.4304/jcp.6.8.1715-1722
- Nerem, R.: 1995, Measuring global mean sea level variations using TOPEX/POSEIDON altimeter data. *J. Geophys. Res. Atmos.*, 100, C12, 25135–25151. DOI: 10.1029/954JC02303
- Oropeza, V. and Sacchi, M.: 2011, Simultaneous seismic data denoising and reconstruction via multichannel singular spectrum analysis. *Geophysics*, 76, 3, 25–32. DOI: 10.1190/1.3552706

- Rajesh, R. and Tiwari, R.: 2014, Windowed SSA (Singular Spectral Analysis) for geophysical time series analysis. *J. Geological Resource and Engineering*, 1, 3, 167-173. DOI: 10.17265/2328-2193/2014.03.004
- Rangelova, E., Sideris, M. and Kim, J.: 2012, On the capabilities of the multi-channel singular spectrum method for extracting the main periodic and non-periodic variability from weekly GRACE data. *J. Geodyn.*, 54, 64-78. DOI: 10.1016/j.jog.2011.10.006
- Robertson, A. and Mechoso, C.: 1998, Interannual and decadal cycles in river flows of southeastern South America. *J. Climate*, 11, 10, 2570-2581. DOI: 10.1175/1520-0442(1998)011<2570:IADCIR>2.0.CO;2
- Schoellhamer, D.: 2001, Singular spectrum analysis for time series with missing data. *Geophys. Res. Lett.*, 28, 16, 3187-3190. DOI: 10.1029/2000GL012698
- Shen, Y., Li, W., Xu, G. and Li, B.: 2014, Spatiotemporal filtering of regional GNSS network's position time series with missing data using principle component analysis. *J. Geodesy*, 88, 1, 1-12. DOI: 10.1007/s00190-013-0663-y
- Shen, Y., Peng, F. and Li, B.: 2015, Improved singular spectrum analysis for time series with missing data. *Nonlinear Proc. Geoph.*, 22, 4, 371-376. DOI: 10.5194/npg-22-371-2015
- Sivapragasam, C., Liong, S. and Pasha, M.: 2001, Rainfall and runoff forecasting with SSA-SVM approach. *J. Hydroinform.*, 3, 3, 141-152.
- Tiwari, R. and Rajesh, R.: 2014, Factorized Hankel optimal singular spectral approach for erratic and noisy seismic signal denoising. *J. Appl. Geophys.*, 111, 95-101. DOI: 10.1016/j.jappgeo.2014.09.019
- Vautard, R. and Ghil, M.: 1989, Singular spectrum analysis in nonlinear dynamics with applications to paleoclimatic time series. *Physica D*, 35, 3, 395-424. DOI: 10.1016/0167-2789(89)90077-8
- Vautard, R., Yiou, P. and Ghil, M.: 1992, Singular-spectrum analysis: A toolkit for short, noisy chaotic signals. *Physica D*, 58, 1-4, 95-126. DOI: 10.1016/0167-2789(92)90103-T
- Vitanov, N., Sakai, K. and Dimitrova, Z.: 2008, SSA, PCA, TDPSC, ACFA: Useful combination of methods for analysis of short and nonstationary time series. *Chaos Soliton Fract.*, 37, 1, 187-202. DOI: 10.1016/j.chaos.2006.08.043
- Watson, P.J.: 2016, Identifying the best performing time series analytics for sea level research. In book *Time Series Analysis and Forecasting*. Springer, 261-278.
- Wu, C. and Chau, K.: 2011, Rainfall-runoff modeling using artificial neural network coupled with singular spectrum analysis. *J. Hydrol.*, 399, 3-4, 394-409.
- Zotov, L., Bizouard, C. and Shum, C.: 2016, A possible interrelation between Earth rotation and climate variability on a decadal time-scale. *J. Geodesy Geodyn.*, 7, 3, 216-222. DOI: 10.1016/j.geog.2016.05.005
- Zotov, L. and Shum, C.: 2010, Multichannel singular spectrum analysis of the gravity field data from grace satellites. *AIP Conference Proceedings*, 1206, 1, 473-479. DOI: 10.1063/1.3292557
- Zotov, L., Sidorenkov, N., Bizouard, C., Shum, C. and Shen, W.: 2017, Multichannel singular spectrum analysis of the axial atmospheric angular momentum. *J. Geodesy Geodyn.*, 8, 6, 433-442. DOI: 10.1016/j.geog.2017.02.010
- Zvonarev, N. and Golyandina, N.: 2017, Iterative algorithms for weighted and unweighted finite-rank time-series approximations. *Stat. Interface*, 10, 1, 5-18. DOI: 10.4310/SII.2017.v10.n1.a1